

Mixing, entrainment and energetics of gravity currents released from two-layer stratified locks

Rui Zhu^{1,2}, Zhiguo He^{1,†} and Eckart Meiburg^{2,†}

¹Ocean College, Zhejiang University, Zhoushan 316021, PR China

²Department of Mechanical Engineering, University of California at Santa Barbara, Santa Barbara, CA 93106, USA

(Received 11 March 2022; revised 31 December 2022; accepted 15 February 2023)

We conduct three-dimensional direct numerical simulations to investigate the mixing, entrainment and energy budgets of gravity currents emerging from two-layer stratified locks. Depending on the density and layer thickness ratios, we find that either the upper layer or lower layer fluid can propagate faster, and that the density structure of the overall gravity current can range from strongly stratified to near-complete mixing. We furthermore observe that intermediate values of the density ratio can maximise mixing between the gravity current layers. Based on the vorticity budget, we propose a theoretical model for predicting the overall gravity current height, along with the front velocity of the two layers, for situations in which the lower layer moves faster than the upper layer. The model identifies the role of the height and thickness ratios in determining the velocity structure of the current, and it clarifies the dynamics of the ambient counter-current. A detailed analysis of the energy budget quantifies the conversion of potential into kinetic energy as a function of the governing parameters, along with the energy transfer between the different layers of the gravity current and the ambient fluid. Depending on the values of the density and layer thickness ratios, we find that the lower lock layer can gain or lose energy, whereas the upper layer always loses energy.

Key words: gravity currents, ocean processes, stratified flows

1. Introduction

Gravity currents are predominantly horizontal flows of one fluid within another, driven by the density difference between these two fluids (Simpson 1982; Härtel, Meiburg & Necker 2000; He *et al.* 2017; Zhao *et al.* 2018). They play an important role in numerous natural

† Email addresses for correspondence: hezhi guo@zju.edu.cn, meiburg@engineering.ucsb.edu

© The Author(s), 2023. Published by Cambridge University Press. This is an Open Access article, distributed under the terms of the Creative Commons Attribution licence (<http://creativecommons.org/licenses/by/4.0/>), which permits unrestricted re-use, distribution and reproduction, provided the original article is properly cited.

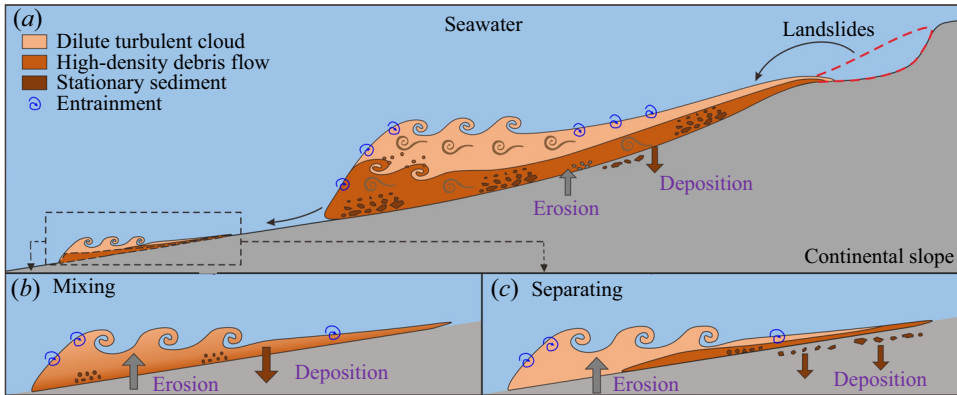


Figure 1. Illustration of the various flow transformations for a subaqueous debris flow. (a) A two-layer subaqueous debris flow. (b) The two layers mix to form a homogeneous current. (c) The upper layer separates from the lower layer.

and industrial processes, and they frequently give rise to strong internal stratification that can take the form of a two-layer structure, with a dilute, less-dense upper layer riding above a dense bottom layer (Cantero *et al.* 2009; Meiburg & Kneller 2010; Sequeiros *et al.* 2010; Kneller *et al.* 2016). An example in this regard concerns subaqueous debris flows with an upper, low-density turbidity cloud (Hampton 1972), as shown in figure 1(a). Depending on the overall flow conditions, the two layers can mix to form a homogeneous current (figure 1b) or they can separate from each other and propagate more or less independently (figure 1c). At present, the mixing and entrainment dynamics giving rise to such flow transformations, along with their energetics, are not well understood.

Several previous studies have analysed various aspects of gravity currents released from two-layer stratified locks that propagate into homogeneous ambient (Gladstone *et al.* 2004; Dai 2017; Wu & Dai 2020; Zemach & Ungarish 2020; He *et al.* 2021). Gladstone *et al.* (2004) found that the propagation of such two-layer lock–release currents is governed by the density and height ratios of the layers. Although weakly stratified layers tend to mix with each other and form homogeneous currents, for strong stratification the lower layer runs ahead of, or is overtaken by, the upper layer and strong stratification persists. Dai (2017) and Wu & Dai (2020) tracked the temporal evolution of the Froude number during the self-similar and slumping phases, respectively. Zemach & Ungarish (2020) presented a self-contained shallow-water model to predict the propagation of inertial gravity currents released from two-layer stratified locks, including channels with non-rectangular cross-sections (Zemach & Ungarish 2021). He *et al.* (2021) proposed a quantitative model for predicting the time required for complete mixing during the propagation of weakly stratified two-layer gravity currents, which has been validated by corresponding experimental data. Although the mixing, entrainment and energy budgets in single-layer gravity currents have been investigated in quite some detail, their counterparts for two-layer currents have received comparatively scant attention.

Numerous previous studies indicate that the entrainment between a single-layer gravity current and the ambient fluid greatly influences the dynamics of the gravity current. Based on experiments employing a colour pH indicator, Hallworth *et al.* (1996) suggested that entrainment occurs mostly during the self-similar phase. On the other hand, Hacker, Linden & Dalziel (1996) applied a light attenuation technique to provide evidence that entrainment occurs during all flow stages. Fragoso, Patterson & Wettlaufer (2013)

forwarded energy arguments in support of entrainment throughout the slumping regime. Samasiri & Woods (2015) and Sher & Woods (2015) used light attenuation techniques to visualise the entrainment process between the current and the ambient fluid. Using experiments and large eddy simulations, Nogueira *et al.* (2014), Samasiri & Woods (2015), Sher & Woods (2015), Ottolenghi *et al.* (2016) as well as Balasubramanian & Zhong (2018) introduced a variety of entrainment parameters and analysed their dependence on the initial excess density, bed roughness, and aspect ratio. Nogueira *et al.* (2014) and Ottolenghi *et al.* (2016) established the influence on the bulk entrainment parameter by the Reynolds, Froude and Richardson numbers. Ross, Dalziel & Linden (2006) and Adduce, Sciortino & Proietti (2012) proposed a parametrisation for entrainment in shallow-water simulations of lock–release gravity currents. Bhaganagar (2017) focused on turbulence generation to demonstrate that substantial entrainment occurs in the head region, and during the early flow stages. The present investigation aims to provide corresponding entrainment information for two-layer gravity currents.

Quantifying the temporal evolution of the energy budget can provide substantial insight into the dynamics of gravity currents. Although this information is frequently difficult to obtain experimentally (Ooi, Constantinescu & Weber 2009), it can usually be acquired in a relatively straightforward fashion from numerical simulations. Winters *et al.* (1995) established a conceptual framework for analysing the energetics of density-stratified Boussinesq flows, which analyses changes in the available potential energy to investigate irreversible diapycnal mixing. Necker *et al.* (2002, 2005) analysed dissipative energy losses in particle-driven gravity currents, with a focus on separating effects at the macroscopic scale from those related to the Stokes flow around individual particles. Ooi *et al.* (2009) focused on the energy budget of lock–exchange compositional gravity currents, whereas Espath *et al.* (2014) compared the energy budgets of two- and three-dimensional simulations. They observed that the dissipation associated with the macroscopic convective fluid motion is smaller in two- than in three-dimensional flows. In the present investigation, we focus on the evolution of the energy budget for two-layer gravity currents, with the quantification of the energy transfer between the upper and lower layers as a key goal.

The present paper is organised as follows. The computational model PARTIES (Biegert, Vowinckel & Meiburg 2017a; Biegert *et al.* 2017b) for performing three-dimensional direct numerical simulations of two-layer gravity currents is presented in § 2, including the governing equations, dimensionless parameters, numerical method and validation results. In § 3, qualitative and quantitative results for the dynamics of two-layer gravity currents are presented and discussed. The key findings of the investigation are summarised in § 4.

2. Problem set-up and computational approach

We generate the desired gravity currents by means of a lock–release process, as sketched in figure 2. The entire flow domain has length \hat{L} , height \hat{H} and width \hat{W} , where $\hat{\cdot}$ indicates a dimensional quantity. The lock of length \hat{L}_s contains two fluid layers of height \hat{h}_U and \hat{h}_L , respectively, with densities $\hat{\rho}_U$ and $\hat{\rho}_L$. The ambient fluid has density $\hat{\rho}_0$. Upon removal of the gate, the two lock fluids form initially distinct gravity currents whose interaction with each other and with the ambient fluid are the subject of the present investigation. Here we consider only full-depth lock–release flows, and we choose \hat{L}_s sufficiently large so that its precise value does not affect the dynamics of the evolving gravity currents.

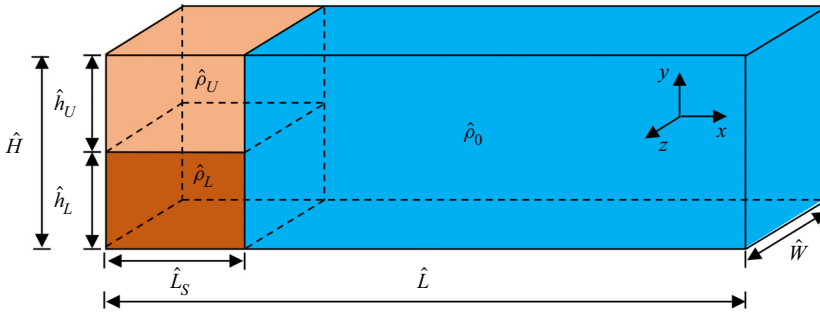


Figure 2. Initial set-up: the lock contains two layers of different densities, whereas the ambient consists of lighter fluid. Upon removal of the gate, the two layers form gravity currents.

2.1. Governing equations

The computational approach follows the earlier work by Necker *et al.* (2002) and Nasr-Azadani & Meiburg (2011). It employs the incompressible Navier–Stokes equations in the Boussinesq approximation, and evolves the concentration fields associated with the upper and lower lock fluids via a pair of advection–diffusion equations:

$$\frac{\partial \hat{u}_j}{\partial \hat{x}_j} = 0, \tag{2.1}$$

$$\frac{\partial \hat{u}_i}{\partial \hat{t}} + \hat{u}_j \frac{\partial \hat{u}_i}{\partial \hat{x}_j} = -\frac{1}{\hat{\rho}_0} \frac{\partial \hat{p}}{\partial \hat{x}_i} + \hat{\nu} \frac{\partial^2 \hat{u}_i}{\partial \hat{x}_j \partial \hat{x}_j} + \frac{\hat{\rho} - \hat{\rho}_0}{\hat{\rho}_0} \hat{g} e_i^g, \tag{2.2}$$

$$\frac{\partial \hat{c}_U}{\partial \hat{t}} + \hat{u}_j \frac{\partial \hat{c}_U}{\partial \hat{x}_j} = \hat{\kappa} \frac{\partial^2 \hat{c}_U}{\partial \hat{x}_j \partial \hat{x}_j}, \tag{2.3}$$

$$\frac{\partial \hat{c}_L}{\partial \hat{t}} + \hat{u}_j \frac{\partial \hat{c}_L}{\partial \hat{x}_j} = \hat{\kappa} \frac{\partial^2 \hat{c}_L}{\partial \hat{x}_j \partial \hat{x}_j}. \tag{2.4}$$

Here \hat{u} denotes the fluid velocity, with the subscripts i and j indicating the x , y or z direction, respectively, $\hat{\rho}$ represents the local density, \hat{t} is time, $\hat{\nu}$ denotes the kinematic viscosity, \hat{g} indicates the gravitational acceleration, with $e^g = (0, -1, 0)$ being the unit vector in the direction of gravity. Even though the concentration fields \hat{c}_U and \hat{c}_L of the upper and lower layer have identical diffusivity $\hat{\kappa}$, we choose to track them separately in order to obtain detailed information about the mixing and energy exchange between the two evolving gravity currents. We assume a linear density–concentration relationship of the form

$$\hat{\rho} = \hat{\rho}_0 + \alpha(\hat{c}_U + \hat{c}_L), \tag{2.5}$$

where α denotes the proportionality factor between concentration and fluid density. The initial upper and lower fluid densities are

$$\hat{\rho}_U = \hat{\rho}_0 + \alpha \hat{c}_{U,init}, \tag{2.6}$$

$$\hat{\rho}_L = \hat{\rho}_0 + \alpha \hat{c}_{L,init}, \tag{2.7}$$

where $\hat{c}_{U,init}$ and $\hat{c}_{L,init}$ indicate the respective initial concentrations.

Two-layer stratified gravity currents

We non-dimensionalise the governing equations by introducing characteristic scales of the form

$$x = \frac{2\hat{x}}{\hat{h}_U}, \quad u = \frac{\hat{u}}{\hat{u}_b}, \quad t = \frac{2\hat{t}\hat{u}_b}{\hat{h}_U}, \quad p = \frac{\hat{p}}{\hat{\rho}_0\hat{u}_b^2}, \quad c_U = \frac{\hat{c}_U}{\hat{c}_{U,init}}, \quad c_L = \frac{\hat{c}_L}{\hat{c}_{L,init}}, \quad (2.8a-f)$$

where $\hat{u}_b = \sqrt{\hat{g}\hat{h}_U(\hat{\rho}_U - \hat{\rho}_0)/2\hat{\rho}_0}$ indicates the buoyancy velocity. The non-dimensional equations thus take the form

$$\frac{\partial u_j}{\partial x_j} = 0, \quad (2.9)$$

$$\frac{\partial u_i}{\partial t} + u_j \frac{\partial u_i}{\partial x_j} = -\frac{\partial p}{\partial x_i} + \frac{1}{Re} \frac{\partial^2 u_i}{\partial x_j \partial x_j} + (c_U + R_\rho c_L) e_i^g, \quad (2.10)$$

$$\frac{\partial c_U}{\partial t} + u_j \frac{\partial c_U}{\partial x_j} = \frac{1}{Pe} \frac{\partial^2 c_U}{\partial x_j \partial x_j}, \quad (2.11)$$

$$\frac{\partial c_L}{\partial t} + u_j \frac{\partial c_L}{\partial x_j} = \frac{1}{Pe} \frac{\partial^2 c_L}{\partial x_j \partial x_j}. \quad (2.12)$$

As governing dimensionless parameters, we obtain the Reynolds number Re ,

$$Re = \frac{\hat{u}_b \hat{h}_U}{2\hat{\nu}}, \quad (2.13)$$

and the Péclet number Pe ,

$$Pe = \frac{\hat{u}_b \hat{h}_U}{2\hat{k}}. \quad (2.14)$$

Here Re and Pe are related by the Schmidt number Sc as

$$Sc = \frac{Pe}{Re} = \frac{\hat{\nu}}{\hat{k}}. \quad (2.15)$$

Compared with single-layer gravity currents, two new dimensionless parameters characterise the initial stratification of the two-layer case. They are the ratio R_ρ of the density differences between the lock fluids and the ambient,

$$R_\rho = \frac{\hat{\rho}_L - \hat{\rho}_0}{\hat{\rho}_U - \hat{\rho}_0}, \quad (2.16)$$

as well as the ratio R_h of the layer thicknesses,

$$R_h = \frac{\hat{h}_L}{\hat{h}_U}. \quad (2.17)$$

We remark that, although previously we have defined the Reynolds number with the density of the upper dense layer, we could alternatively have employed the density of the lower layer, which would provide a corresponding Reynolds number Re_L :

$$Re_L = \sqrt{R_\rho} Re. \quad (2.18)$$

For clarity, in the following we provide both Reynolds numbers for the simulations to be discussed.

Run	h_U	h_L	R_h	R_ρ	Re	Re_L
1	2	0	0.00	—	7000	—
2	2	0.2	0.10	1.25	7000	7826
3	2	0.2	0.10	2.00	7000	9899
4	2	0.2	0.10	4.00	7000	14 000
5	2	0.5	0.25	1.25	7000	7826
6	2	0.5	0.25	2.00	7000	9899
7	2	0.5	0.25	4.00	7000	14 000
8	2	2.0	1.00	1.00	7000	7000
9	2	2.0	1.00	1.25	7000	7826
10	2	2.0	1.00	2.00	7000	9899
11	2	2.0	1.00	4.00	7000	14 000
12	2	3.0	1.50	1.25	7000	7826
13	2	3.0	1.50	2.00	7000	9899
14	2	3.0	1.50	4.00	7000	14 000

Table 1. Simulation cases and parameters.

Hereafter, we set $Sc = 1$ for simplicity, as its effect on the flow is not a main focus on the present investigation (Necker *et al.* 2005). The evolving gravity current flow is thus fully characterised by Re , R_ρ and R_h . In the following, we generally take Re sufficiently large so that the influence of its exact value on the global flow properties is small. Thus, our primary goal is to investigate the influence of R_ρ and R_h on the interaction among the lock and ambient fluids.

2.2. Initial and boundary conditions

In all cases, the domain is a rectangular tank with $L = 50$ and $W = 1$. The length of the lock is $L_s = 10$. The initial non-dimensional concentrations within the upper and lower layer fluids are $c_{U,init} = 1$ and $c_{L,init} = 1$, respectively, and the concentration within the ambient fluid is zero. The parameters of the simulations are presented in table 1.

We impose no-slip boundary conditions along the bottom and side boundaries ($y = 0$ and $x = 0, L$), whereas a free-slip condition holds along the top boundary ($y = H$). For the concentration fields, no-flux conditions are implemented along all boundaries in the x and y directions. Periodic boundaries are assumed in the z direction for both velocity and concentration.

2.3. Numerical method

We employ our in-house incompressible Navier–Stokes solver PARTIES (Biegert *et al.* 2017a,b) to conduct all simulations. It uses second-order central finite differences to discretise the viscous and diffusive terms in the momentum and transport equations, along with a second-order upwind scheme for the advection terms. The viscous and diffusive terms are solved implicitly to ensure that these terms do not result in any restrictions on the time step (Nasr-Azadani & Meiburg 2011). Time integration is performed by means of a third-order low-storage Runge–Kutta method. The pressure-projection method is implemented, based on a direct fast Fourier transform (FFT) solver for the resulting Poisson equation at each Runge–Kutta substep. We employ a uniform mesh size $l = 0.01$ in all three directions, which satisfies the requirement for adequate resolution

Two-layer stratified gravity currents

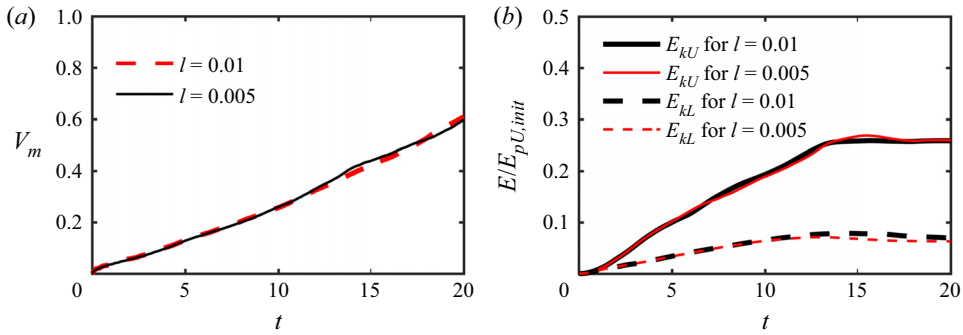


Figure 3. Time history of (a) the normalised mixed volume V_m and (b) the kinetic energy of the upper layer E_{pU} and lower layer E_{pL} for the simulations with $l = 0.01$ and $l = 0.005$ with $R_\rho = 2$ and $R_h = 0.25$.

(Härtel *et al.* 1997, 2000; Birman, Martin & Meiburg 2005) as

$$l \approx \frac{1}{\sqrt{ReSc}}. \quad (2.19)$$

To demonstrate the convergence of the numerical results, we simulated run 6 for a finer grid with $l = 0.005$. Figure 3 compares the normalised mixed volume V_m and energy terms for the case of $l = 0.01$ with the simulation of $l = 0.005$. These quantities are essentially identical for different grid sizes, which indicates that the results are converged. Details regarding the energy terms and the mixed volume are given in §§ 3.3 and 3.4, respectively.

2.4. Validation of the numerical model

A variety of validation results for PARTIES are presented in earlier work by Nasr-Azadani & Meiburg (2011) and Nasr-Azadani, Hall & Meiburg (2013). In order to demonstrate its suitability for the propagation of two-layer gravity currents, we compare simulation results with corresponding experimental data for a representative two-layer gravity current flow. The experimental set-up, described in He *et al.* (2021), involves a tank of $2\text{ m} \times 0.2\text{ m} \times 0.2\text{ m}$, with a watertight gate 0.1 m from one end. The ambient fluid consists of fresh water, whereas the dense fluid layers in the lock are saline solutions with different densities, cf. table 2. Figure 4(a) compares the experimental ($X_{f,exp}$) and numerical ($X_{f,sim}$) front locations as functions of time. Here, we define the front position as the rightmost location in the flow field with c_U or c_L equal to 0.5. The mean error $E_{er}(t) = \text{mean}(|X_{f,exp}(t) - X_{f,sim}(t)|/X_{f,exp}(t))$ generally is $O(10\%)$, which indicates that the numerical results are in good agreement with the experimental values. The simulation values are slightly larger than the experimental values, which is likely due to the fact that the simulation employs periodic boundaries in the spanwise direction, whereas the experiments have solid walls, which will retard the front propagation. The quantity E_{er} is also evaluated by Ottolenghi *et al.* (2016) in order to assess simulation accuracy. Figure 4(b) compares the horizontal velocity profiles for the experiment and numerical simulation. These profiles are averaged in spanwise and streamwise directions over three lock lengths behind the front of the currents at $t = 23$. The values are normalised by u_{max} and $y_{0.5}$ (Buckee, Kneller & Peakall 2001; Gerber, Diedericks & Basson 2011). Here u_{max} is the maximum velocity of the profile and $y_{0.5}$ is the height at which the velocity is half the maximum value. The experimental and simulation results show acceptable agreement.

	\hat{h}_U (m)	\hat{h}_L (m)	$\hat{\rho}_U$ (kg m ⁻³)	$\hat{\rho}_L$ (kg m ⁻³)	$\hat{\rho}_0$ (kg m ⁻³)
Exp. 1	0.09	0.06	1001.7	1017.6	997.8

Table 2. Experimental parameters for the validation case.

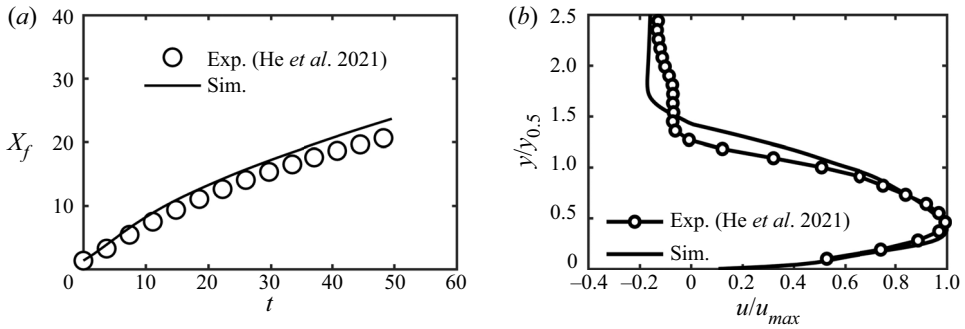


Figure 4. Comparison of experimental and computational (a) front locations and (b) velocity profiles for the validation case in table 2.

3. Results

3.1. Flow regimes

As a representative example, figure 5 shows the evolution of the two-layer gravity current for run 10, with $R_\rho = 2$ and $R_h = 1$. Upon removal of the gate, the lower layer (in blue) develops a distinct head, whereas the upper layer (in red) forms a wedge-like intrusion behind the lower-layer head, cf. figure 5(a). The shear along the interface between the ambient and dense fluids results in the formation of strong Kelvin–Helmholtz vortices. As the current propagates further, we can distinguish a dense head of lower-layer fluid and a more dilute wake containing both upper and lower-layer fluid (figure 5b). These findings are similar to previous experimental observations (Gladstone *et al.* 2004; Dai 2017; Wu & Dai 2020; He *et al.* 2021). The wake region immediately behind the head is characterised by strong turbulent mixing of the lock fluid layers, whereas further back, the two dense layers remain more distinct, as seen in figure 5(c). The mixed region gradually expands, as dense fluid from the head is transported into the turbulent wake (figure 5d).

Figure 6 compares three cases with identical thickness ratio $R_h = 1$, but different density ratios R_ρ , i.e. runs 2–4. For a moderate density ratio $R_\rho = 1.25$, the fronts emerging from the upper and lower lock layers initially have comparable velocities, and significant mixing occurs near the head of the current (figure 6a). For larger density ratios ($R_\rho = 2$ and 4), on the other hand, the lower layer front propagates much faster than the upper layer front, so that the current head consists nearly entirely of lower-layer fluid, which was also observed by Wu & Dai (2020). The upper, red layer advances much more slowly, giving rise to much weaker Kelvin–Helmholtz vortices along its interface with the ambient fluid, along with much less-vigorous mixing (figure 6b,c).

Figure 7 focuses on flows with constant density ratio $R_\rho = 4$ and different thickness ratios R_h , i.e. runs 1, 4, 7 and 14. For $R_h = 0$, the lower layer is absent and the upper layer propagates as a single-layer gravity current (figure 7a). For the small but finite value $R_h = 0.1$, the thin lower layer lubricates the much thicker upper layer, so that the upper layer advances faster than for $R_h = 0$. Note that the head of the upper layer outruns the

Two-layer stratified gravity currents

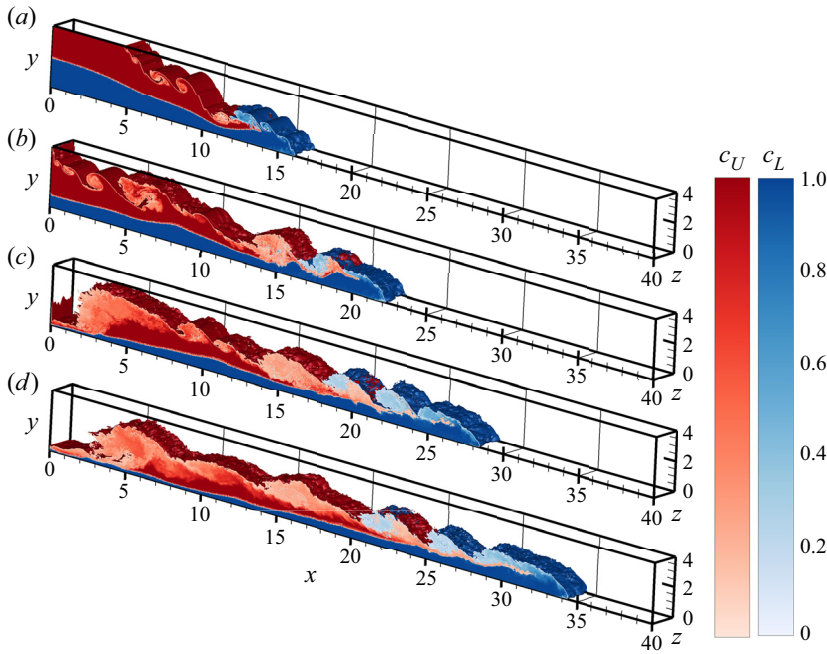


Figure 5. Instantaneous concentration isosurfaces $c_U = 0.1$ and $c_L = 0.1$, along with the concentration field for both upper and lower layers for run 10 ($R_\rho = 2$ and $R_h = 1$), (a) $t = 5$, (b) $t = 10$, (c) $t = 15$ and (d) $t = 20$.

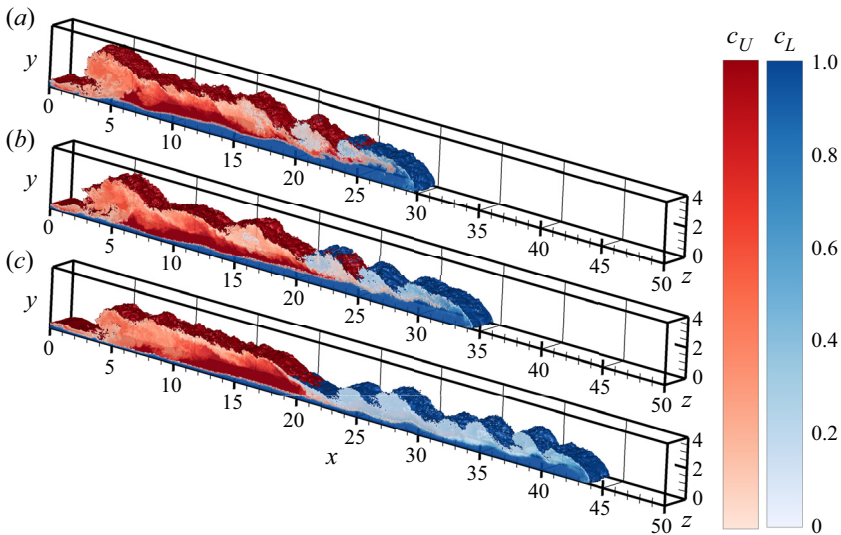


Figure 6. Instantaneous concentration isosurfaces $c_U = 0.1$ and $c_L = 0.1$, along with the concentration field for both upper and lower layers at $t = 20$ for $R_h = 1$: (a) run 2 for $R_\rho = 1.25$, (b) run 3 for $R_\rho = 2$ and (c) run 4 for $R_\rho = 4$.

lower layer front (figure 7b). A similar trend was also observed by Gladstone *et al.* (2004), Dai (2017), Wu & Dai (2020) and He *et al.* (2021). The conditions under which this can occur are discussed in more detail in the following section. For $R_h = 0.25$ the two front velocities are similar, and significant mixing occurs in the head region (figure 7c).

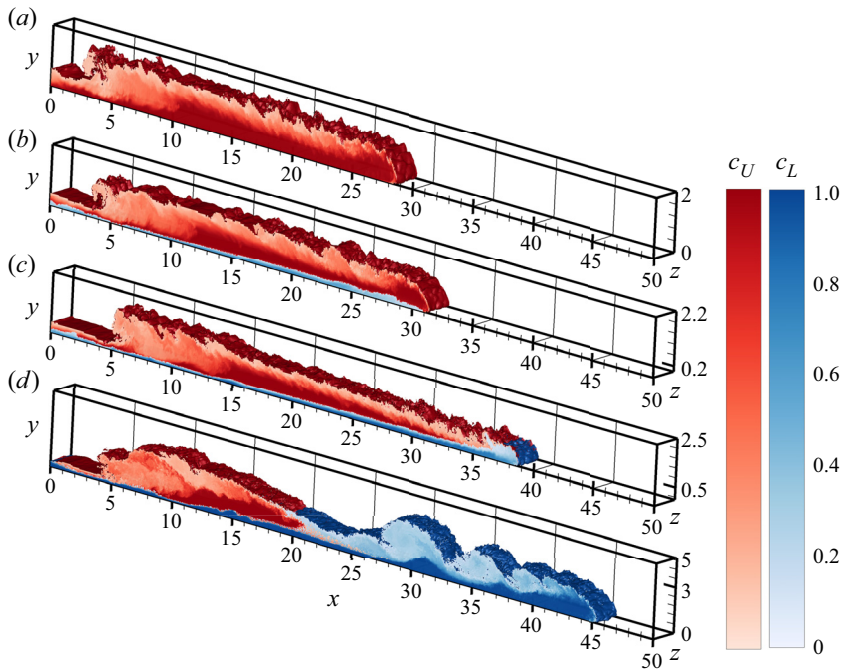


Figure 7. Instantaneous concentration isosurfaces $c_U = 0.1$ and $c_L = 0.1$, along with the concentration field for both upper and lower layers for $R_\rho = 4$: (a) run 1 for $R_h = 0$ at $t = 30$, (b) run 4 for $R_h = 0.1$ at $t = 30$, (c) run 7 for $R_h = 0.25$ at $t = 30$ and (d) run 14 for $R_h = 1.5$ at $t = 18$.

For $R_h = 1.5$, on the other hand, the lower layer front propagates faster than the upper layer, so that the current head consists mostly of lower-layer fluid (figure 7d). We now proceed to a more quantitative analysis of the dependence of the front velocity, energy budget, mixing and entrainment on R_ρ and R_h .

3.2. Front velocity

We define the front positions $X_{f,U}$ and $X_{f,L}$ of the upper and lower layers as the rightmost locations in the flow field with $c_U = 0.5$ and $c_L = 0.5$, respectively. This large threshold value is selected so that small amounts of fluid from one layer that have become entrained into the other layer do not result in a spurious front location value. The respective front velocities are then obtained as $U_{f,U} = dX_{f,U}/dt$ and $U_{f,L} = dX_{f,L}/dt$.

Figure 8 displays the front velocities as functions of time for runs 4 and 10. For single-layer currents, it is well known that they proceed through a short acceleration phase into a constant-velocity slumping phase that can last for an extended period of time and typically extends over approximately four lock lengths (Rottman & Simpson 1983). Figure 8 indicates that the two-layer gravity currents of the present study stay within the slumping phase throughout the entire simulation, as their front velocity, which is determined by the rate at which the more advanced layer propagates, remains approximately constant. The slower layer, on the other hand, displays a highly unsteady front velocity, as it is strongly affected by the turbulent wake of the current head.

Figure 8(a) shows that for run 10 the velocity of the lower, faster layer remains approximately constant, whereas that of the upper, slower layer fluctuates with time. Figure 8(b), on the other hand, indicates that for run 4 the upper, faster layer propagates at

Two-layer stratified gravity currents

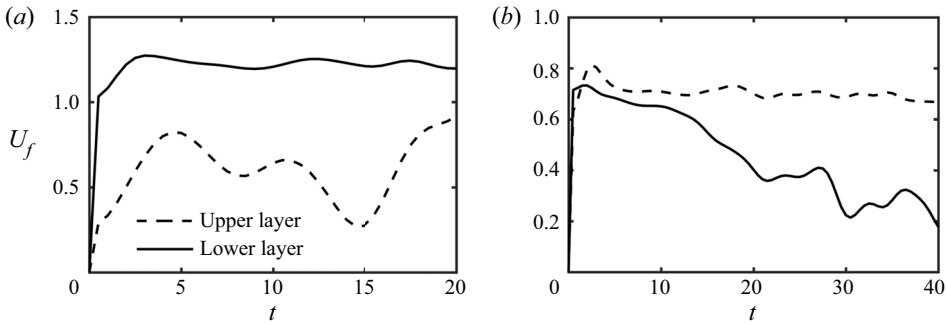


Figure 8. Upper and lower layer front velocities U_f as a function of time for different density ratios R_ρ and height ratios R_h . (a) The lower layer moves ahead of the upper layer for run 10 ($R_\rho = 2$, $R_h = 1$). (b) The upper layer overtakes the lower layer for run 4 ($R_\rho = 4$, $R_h = 0.1$).

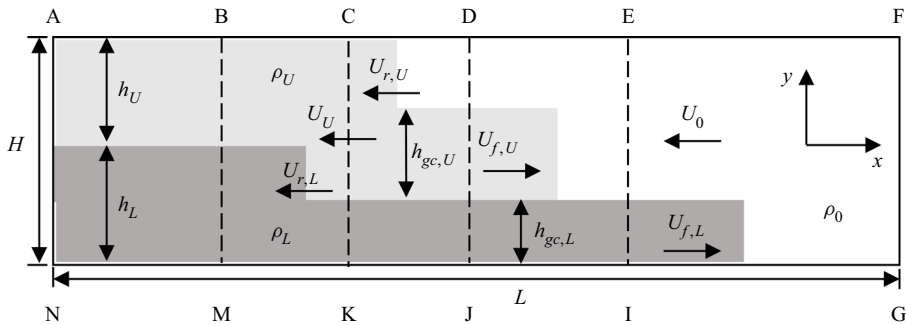


Figure 9. Model flow for predicting the upper and lower layer front velocities $U_{f,U}$ and $U_{f,L}$.

an approximately constant velocity, whereas the lower, slower layer decelerates and does not develop a quasi-steady front velocity.

We now proceed to employ the conservation of mass and vorticity to develop a circulation model for the case of a faster lower layer, by building on the earlier work by Borden & Meiburg (2013). Figure 9 presents the schematic flow field. After removal of the gate, the lower-layer lock fluid forms a right-propagating gravity current front of height $h_{gc,L}$ with velocity $U_{f,L}$. Far behind the front, a discontinuity propagates with velocity $U_{r,L}$ to the left. Simultaneously, a counter current of velocity U_0 emerges in the ambient fluid above the gravity current. Subsequently, the upper-layer lock fluid also forms a gravity current of height $h_{gc,U}$ and velocity $U_{f,U}$ that propagates on top of the lower layer fluid. The flow is fully described by eight dimensionless unknowns: $U_{f,L}$, U_0 , $U_{r,L}$, $U_{f,U}$, U_U , $U_{r,U}$, $h_{gc,L}$ and $h_{gc,U}$.

Within control volume EFGI, and in the reference frame moving with the lower layer fluid, mass conservation for the ambient fluid gives

$$U_{f,L}h_{gc,L} = U_0(H - h_{gc,L}). \quad (3.1)$$

Because no vorticity enters the control volume, and vorticity flows out of the control volume confined to a thin vortex sheet between the lower layer and the ambient fluid, the vorticity flux is equal to the vortex sheet strength $U_0 + U_{f,L}$, multiplied by the sheet's principal velocity $(U_0 + U_{f,L})/2$. The vorticity conservation at the interface gives

$$R_\rho h_{gc,L} = \frac{1}{2}(U_0 + U_{f,L})^2. \quad (3.2)$$

In control volume DEIJ, in the reference frame moving with the upper layer, the mass conservation for the ambient fluid, along with vorticity conservation at the upper surface of the upper layer can be written as

$$U_0(H - h_{gc,L}) + U_{f,U}h_{gc,U} = U_{r,U}(H - h_{gc,L} - h_{gc,U}), \quad (3.3)$$

$$h_{gc,U} = \frac{1}{2}(U_{r,U} + U_{f,U})^2. \quad (3.4)$$

For control volume CDJK, in the frame moving with the disturbance of the upper layer, mass and vorticity conservation yields

$$U_{r,U}(H - h_{gc,L} - h_{gc,U}) = U_U(H - h_{gc,L}) + U_{f,U}h_{gc,U}, \quad (3.5)$$

$$H - h_{gc,L} - h_{gc,U} = \frac{1}{2}(U_{f,U} + U_{r,U})^2. \quad (3.6)$$

Similarly in control volume BCKM, in the frame moving with the disturbance of the lower layer, mass and vorticity conservation leads to

$$U_U(H - h_{gc,L}) = U_{r,L}(H - h_{gc,L} - h_U), \quad (3.7)$$

$$(R_\rho - 1)(h_L - h_{gc}^L) = (U_{f,L} + U_U) \left(\frac{U_{f,L} - U_U}{2} + U_{r,L} \right). \quad (3.8)$$

By combining these equations, we obtain for the dimensionless front velocity and height of the upper and lower layer gravity currents

$$U_{f,L} = \frac{(2 + 2R_h - h_{gc,L})\sqrt{2R_\rho h_{gc,L}}}{2(1 + R_h)}, \quad U_{f,U} = \frac{\sqrt{2h_{gc,U}}}{2} - \frac{h_{gc,L}\sqrt{2R_\rho h_{gc,L}}}{2(1 + R_h)}. \quad (3.9a,b)$$

$$h_{gc,L} = \frac{R_h(1 + R_h)(3R_\rho - 2) - R_h\sqrt{R_\rho(1 + R_h)[R_\rho(9 + R_h) - 8]}}{R_h(2R_\rho - 1) - 1}, \quad (3.10)$$

$$h_{gc,U} = \frac{(1 + R_h)(R_\rho R_h - 2) + R_h\sqrt{R_\rho(1 + R_h)[R_\rho(9 + R_h) - 8]}}{2R_h(2R_\rho - 1) - 2}. \quad (3.11)$$

Figure 10 compares simulation results for the upper and lower layer front velocities with corresponding experimental data of He *et al.* (2021), and with predictions by (3.9a,b). We observe generally good agreement, as reflected by the coefficient of determination $R^2 = 0.814$. We note that for $R_h = 0$, i.e. in the absence of a lower layer, (3.9a,b) yields the correct solution $U_{f,U} = \sqrt{2}/2$ for single-layer currents (Borden & Meiburg 2013). In the limit of $R_\rho = 1$, the two layers have identical density and form a single-layer current with $U_{f,U} = \sqrt{2(R_h + 1)}/2$, which is again consistent with Borden & Meiburg (2013). If we mix the two lock layers initially so as to form a homogeneous lock fluid, a single-layer current forms with front velocity $\sqrt{2(1 + R_h R_\rho)}/2$, based on the circulation model and the half-depth assumption. Corresponding data, indicated in figure 10, suggest that the front velocity of a two-layer current is quantitatively similar to that of an equivalent, initially well-mixed single-layer current.

Figure 11 shows predictions for the front velocities by (3.9a,b) as functions of R_h , for different values of R_ρ . Here $U_{f,L}$ increases with R_h and R_ρ . When R_ρ is small, $U_{f,U}$ increases slightly for all R_h . For larger R_ρ , on the other hand, $U_{f,U}$ decreases for all R_h values. This illustrates the competing effects of increased potential energy conversion and larger ambient counter-current velocity. We furthermore note that as R_ρ increases, a smaller value of R_h will suffice to ensure that the lower layer propagates faster than the upper layer.

Two-layer stratified gravity currents

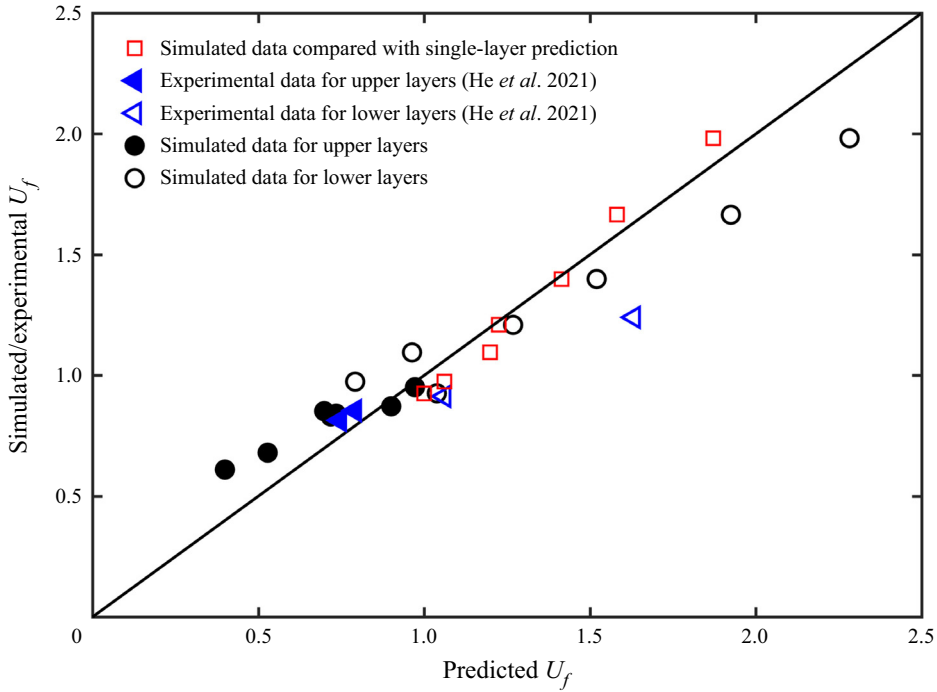


Figure 10. The black and blue symbols compare simulation results for the upper and lower layer front velocities, along with experimental data by He *et al.* (2021), to corresponding predictions by (3.9a,b). The experimental data are for cases with $R_\rho = 5$ for $R_h = 0.34$ and 0.13. For all cases the lower layer moves faster than the upper layer. Also shown by red symbols are the corresponding simulation results for single-layer currents and predictions by $\sqrt{2(1 + R_h R_\rho)}/2$.

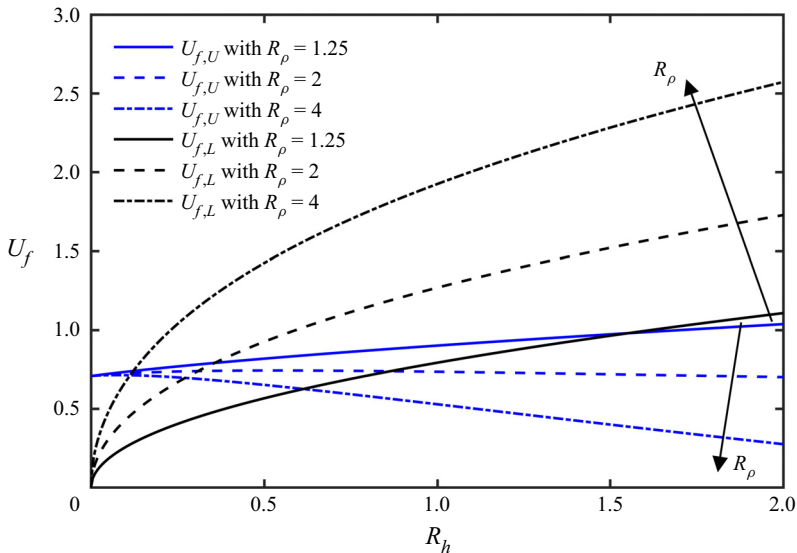


Figure 11. Predictions by (3.9a,b) for the upper and lower layer front velocities, $U_{f,U}$ and $U_{f,L}$, as functions of R_h , for different values of R_ρ .

3.3. Energy budget

The initial potential energy stored in the upper and lower dense fluid layers acts as the source that drives the flow. Once the gate opens, this potential energy is partially converted into kinetic energy, which is subsequently dissipated by the action of viscosity. In the following, we analyse these energy conversion processes quantitatively, along with the energy transfer between the different fluid layers. Towards this end, we define the potential energy E_p , kinetic energy E_k and dissipated energy E_d , respectively, as

$$E_p(t) = \int_0^W \int_0^H \int_0^L (c_U + R_\rho c_L)(y - h_L) \, dx \, dy \, dz, \quad (3.12)$$

$$E_k(t) = \int_0^W \int_0^H \int_0^L \frac{1}{2}(u^2 + v^2 + w^2) \, dx \, dy \, dz, \quad (3.13)$$

$$E_d(t) = \int_0^t \varepsilon(\tau) \, d\tau, \quad (3.14)$$

where u , v and w indicate the fluid velocity components in the x , y and z directions. Note that we focus on the potential energy relative to the situation with ambient fluid everywhere. We choose the initial top boundary of the lower dense layer $y = h_L$ as our reference level, so that the initial potential energy contained in the upper lock layer is constant for all cases. Here ε indicates the instantaneous viscous dissipation rate

$$\varepsilon(t) = \int_0^W \int_0^H \int_0^L \frac{2}{Re} s_{ij} s_{ij} \, dx \, dy \, dz, \quad (3.15)$$

with s_{ij} denoting the rate-of-strain tensor

$$s_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right). \quad (3.16)$$

As the fluid is at rest initially, the overall energy budget of the flow takes the form

$$E_{total} = E_p + E_k + E_d = \text{const.} = E_{p,init}, \quad (3.17)$$

where $E_{p,init}$ represents the initial potential energy. To analyse the time-dependent conversion of potential, kinetic and dissipated energy, and the energy transfer among the different layers, we consider the respective contributions of the individual layers to the overall energy budget:

$$E_k = E_{kU} + E_{kL} + E_{k0}, \quad (3.18)$$

$$E_p = E_{pU} + E_{pL}, \quad (3.19)$$

$$E_d = E_{dU} + E_{dL} + E_{d0}, \quad (3.20)$$

$$\varepsilon = \varepsilon_U + \varepsilon_L + \varepsilon_0, \quad (3.21)$$

Two-layer stratified gravity currents

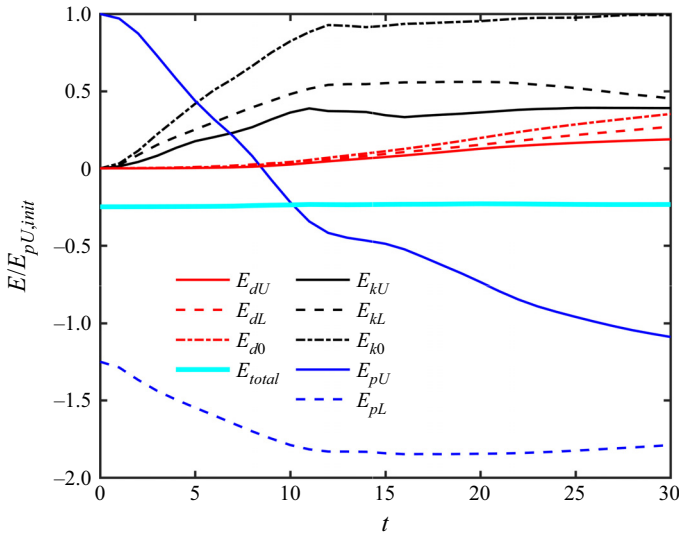


Figure 12. Time history of the various energy budget components for run 9 with $R_\rho = 1.25$ and $R_h = 1$.

where the subscripts U , L and 0 again refer to the upper layer, lower layer and ambient fluid, respectively. These contributions take the form

$$E_{kU}(t) = \int_0^W \int_0^H \int_0^L c_U \frac{1}{2} (u^2 + v^2 + w^2) dx dy dz, \quad (3.22)$$

$$E_{kL}(t) = \int_0^W \int_0^H \int_0^L c_L \frac{1}{2} (u^2 + v^2 + w^2) dx dy dz, \quad (3.23)$$

$$E_{k0}(t) = \int_0^W \int_0^H \int_0^L (1 - c_U - c_L) \frac{1}{2} (u^2 + v^2 + w^2) dx dy dz, \quad (3.24)$$

$$E_{pU}(t) = \int_0^W \int_0^H \int_0^L c_U (y - h_L) dx dy dz, \quad (3.25)$$

$$E_{pL}(t) = \int_0^W \int_0^H \int_0^L R_\rho c_L (y - h_L) dx dy dz, \quad (3.26)$$

$$\varepsilon_U(t) = \int_0^W \int_0^H \int_0^L c_U \frac{2}{Re} s_{ij} s_{ij} dx dy dz, \quad (3.27)$$

$$\varepsilon_L(t) = \int_0^W \int_0^H \int_0^L c_L \frac{2}{Re} s_{ij} s_{ij} dx dy dz, \quad (3.28)$$

$$\varepsilon_0(t) = \int_0^W \int_0^H \int_0^L (1 - c_U - c_L) \frac{2}{Re} s_{ij} s_{ij} dx dy dz. \quad (3.29)$$

Figure 12 shows the time evolution of all energy components for the representative example of run 9 with $R_\rho = 1.25$ and $R_h = 1$. Note that all energy components are normalised by the initial potential energy of the upper dense layer $E_{pU,init}$. The total energy E_{total} varies by around 1.5 % of $E_{pU,init}$ over the course of the simulation, which suggests that energy is conserved with good accuracy throughout the simulation.

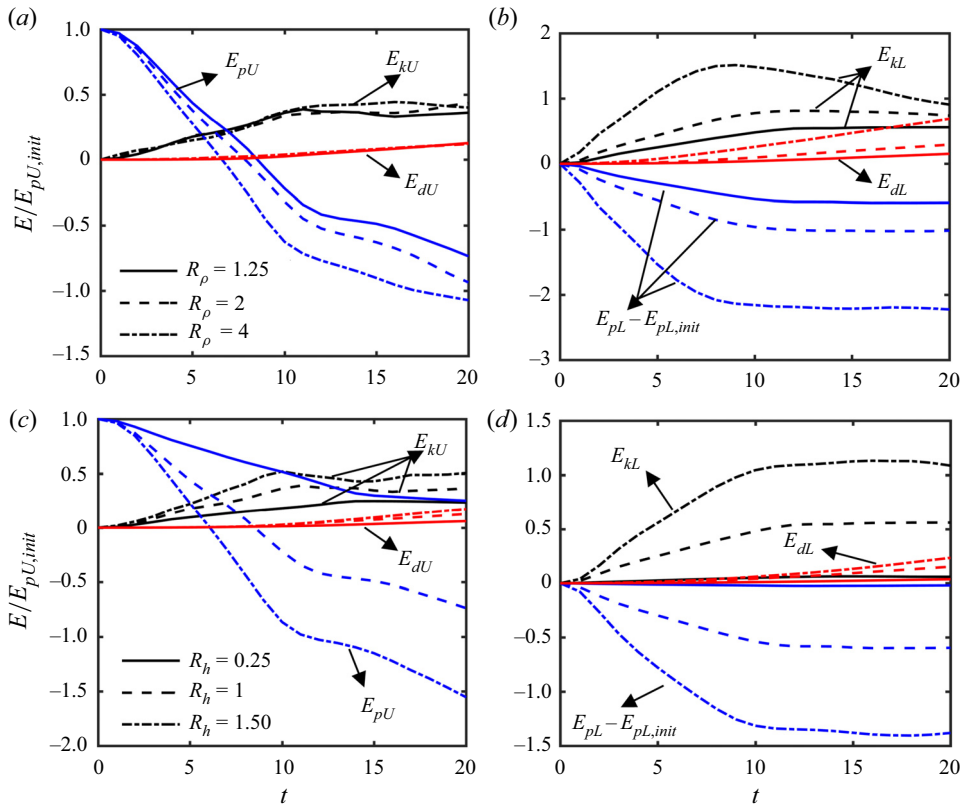


Figure 13. The evolution of the various energy budget components for different R_ρ and R_h : (a) upper layer energy components for $R_h = 1$ and different R_ρ values, i.e. runs 9–11; (b) lower layer energy components for $R_h = 1$ and different R_ρ values; (c) upper layer energy components for $R_\rho = 1.25$ and different R_h , i.e. runs 5, 9 and 12; and (d) lower layer energy components for $R_\rho = 1.25$ and different R_h .

For the first 10 time units the potential energy of the upper and lower dense layers rapidly decreases as the lock fluid collapses, whereas the kinetic energy of all three layers increases. Around $t \approx 10$, the collapse of the lower layer is nearly complete and its potential energy levels off. For late times, its potential energy even increases slightly due to mixing processes that tend to lift up parcels of dense fluid. On the other hand, the upper dense fluid layer continues to lose potential energy. The kinetic energy of all three layers levels off after $t \approx 10$. For late times E_{kU} and E_{kL} gradually decrease due to the effects of dissipation. The dissipated energy grows slowly with time for all three layers.

Figure 13 shows the influence of the density ratio R_ρ and the height ratio R_h on the evolution of the different energy components. Figure 13(a) indicates that the kinetic energy and dissipation of the upper layer do not vary significantly with the lower layer density, even though the upper layer loses potential energy faster in the presence of a denser lower layer that collapses more quickly. Figure 13(b) shows that the energetics of the lower layer are affected more strongly by an increase in R_ρ . As the density of the lower layer increases, it converts potential energy into kinetic energy more quickly, and it dissipates energy at a faster rate. As a larger height ratio R_h also causes the lower layer to propagate faster, R_h has a similar influence on the lower layer energy components as R_ρ , cf. figure 13(d). On the other hand, the dynamics of the upper layer are affected more strongly by R_h than by R_ρ , cf. figure 13(c). Here E_{kU} and E_{dU} increase with R_h , because for larger R_h , the upper layer

Two-layer stratified gravity currents

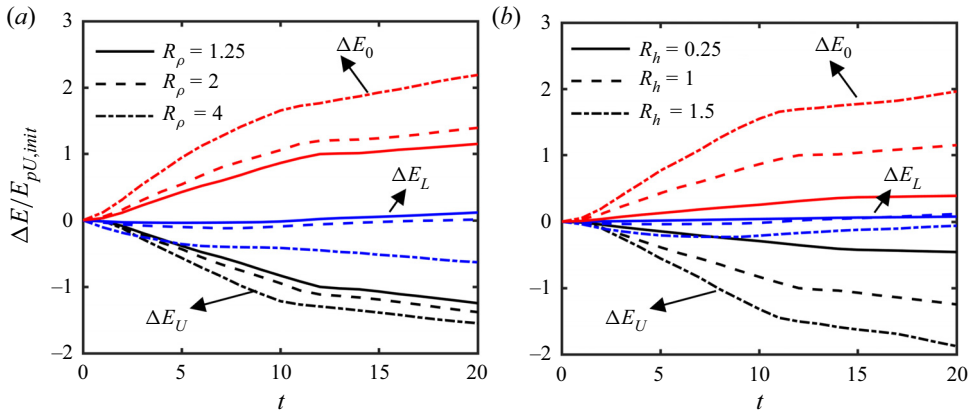


Figure 14. Time history of the variation of the total energy of the upper layer fluid ΔE_U , the lower layer fluid ΔE_L and the ambient fluid ΔE_0 for the cases with (a) $R_h = 1$ and different R_ρ , i.e. runs 9–11, and (b) $R_\rho = 1.25$ and different R_h , i.e. runs 5, 9 and 12.

can descend to a lower vertical level, which enables it to convert more potential energy into kinetic energy.

For a single-layer gravity current, the dense fluid always transfers energy to the ambient fluid, e.g. Necker *et al.* (2002) and He *et al.* (2018). For a two-layer gravity current the situation becomes more complex, as shown in figure 14. In this figure we analyse the change in total energy for each individual fluid layer, ΔE_U , ΔE_L and ΔE_0 , in order to obtain insight into the energy transfer between the different layers. Although the ambient fluid gains energy for all values of R_ρ and R_h , and the upper dense layer always loses energy, the lower dense layer can gain or lose energy, depending on the specific values of the governing dimensionless parameters. This net energy gain of the lower dense layer, as a result of energy transfer from the upper layer, is perhaps somewhat unexpected. To quantify this effect as function of R_ρ and R_h , we focus on the normalised ΔE_U and ΔE_L at the representative time $t = 20$. Figure 15(a) presents ΔE_U as a function of R_h for different values of R_ρ . We find that to a good approximation ΔE_U varies linearly with R_h . Thus, we look for a scaling parameter of the form $R_\rho^{11} R_h$. Figure 15(b) shows that the raw data are well fitted ($R^2 = 0.997$) by a straight line of the form

$$\frac{\Delta E_U}{E_{pU,init}} = -0.2 - 1.1 R_\rho^{0.2} R_h. \quad (3.30)$$

The intersect with the vertical axis at -0.2 indicates that the upper layer loses about 20 % of its initial energy when $R_h = 0$. This suggests that a single-layer gravity current with a small aspect ratio will transfer approximately 20 % of its initial energy to the ambient fluid, as the lock fluid collapses and forms the gravity current. A similar conclusion can also be found from Zhu, He & Meiburg (2021). On the other hand, the energy variation of the upper layer mainly comes from the decrease of the potential energy E_{pU} , which is shown in figure 13. E_{pU} changes approximately linearly with R_h . However, a larger R_ρ results in a faster lower layer, which enhances the loss of E_{pU} indirectly. Equation (3.30) captures the linear and nonlinear dependence of E_{pU} on R_h and R_ρ , respectively, and it reflects that the decrease of E_U depends more strongly on R_h than on R_ρ .

Figure 15(c) shows the normalised ΔE_L as a function of the height ratio R_h for different density ratios R_ρ . As R_h increases, ΔE_L initially increases as well, but it subsequently

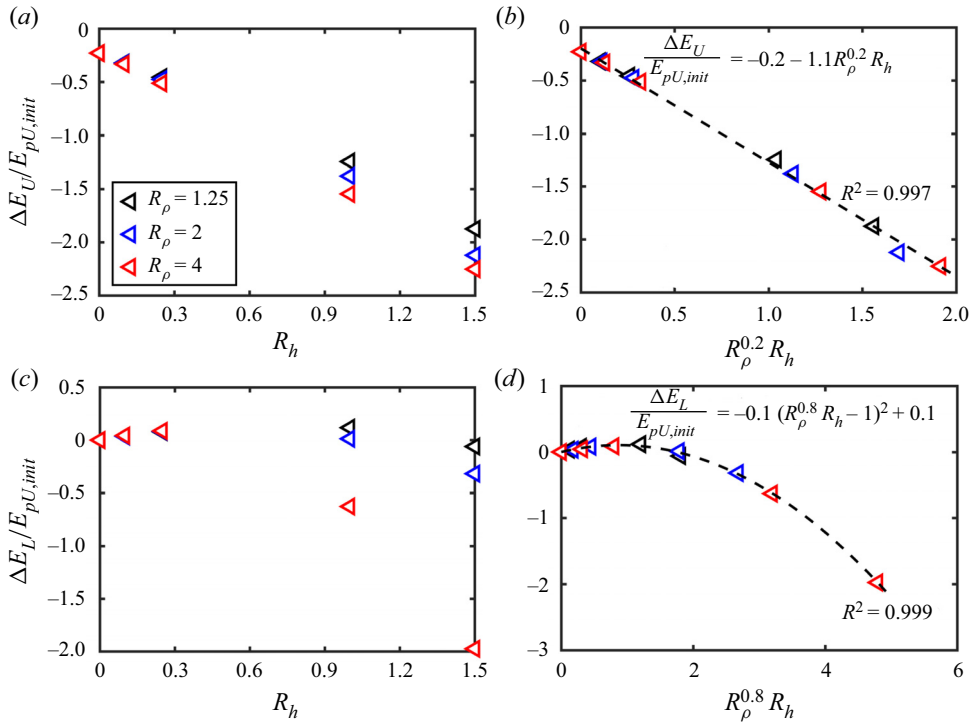


Figure 15. Variation of the total energy of the upper layer fluid ΔE_U , and of the lower layer fluid ΔE_L , for different values of R_ρ and R_h .

reaches a maximum and then decreases. This suggests that for small R_h the lower layer gains more energy from the upper layer than it transfers to the ambient fluid, whereas for large R_h this situation is reversed. We find that the raw data depend parabolically on $R_\rho^{0.8} R_h$, as seen in figure 15(d). A good fit ($R^2 = 0.999$) is achieved by

$$\frac{\Delta E_L}{E_{pU,init}} = -0.1(R_\rho^{0.8} R_h - 1)^2 + 0.1. \tag{3.31}$$

When $R_h = 0$, the lower layer does not exist and ΔE_L is zero. Equation (3.31) also indicates that the lower layer gains a maximum of 10% of the initial upper layer energy, when $R_\rho^{0.8} R_h = 1$. Here $R_\rho^{0.8} R_h$ can be roughly interpreted as representing the mass of the lower layer. A lower layer of larger mass obtains more energy from the upper layer, up to a threshold beyond which the effect of the upper layer on the lower layer starts to wane, and the lower layer transfers more energy to the ambient fluid.

3.4. Mixing of the upper and lower layers

We now proceed to analyse and quantify the mixing dynamics between the upper and lower dense fluid layers, which strongly affects the overall propagation rate of the two-layer current. In order to obtain insight into the Lagrangian mixing behaviour, we mark individual fluid patches by passive tracers of different colours, so that we can easily track them throughout the evolution of the flow, cf. figure 16. Depending on whether the lower dense layer advances faster or more slowly than the upper layer, we observe different mixing patterns.

Two-layer stratified gravity currents

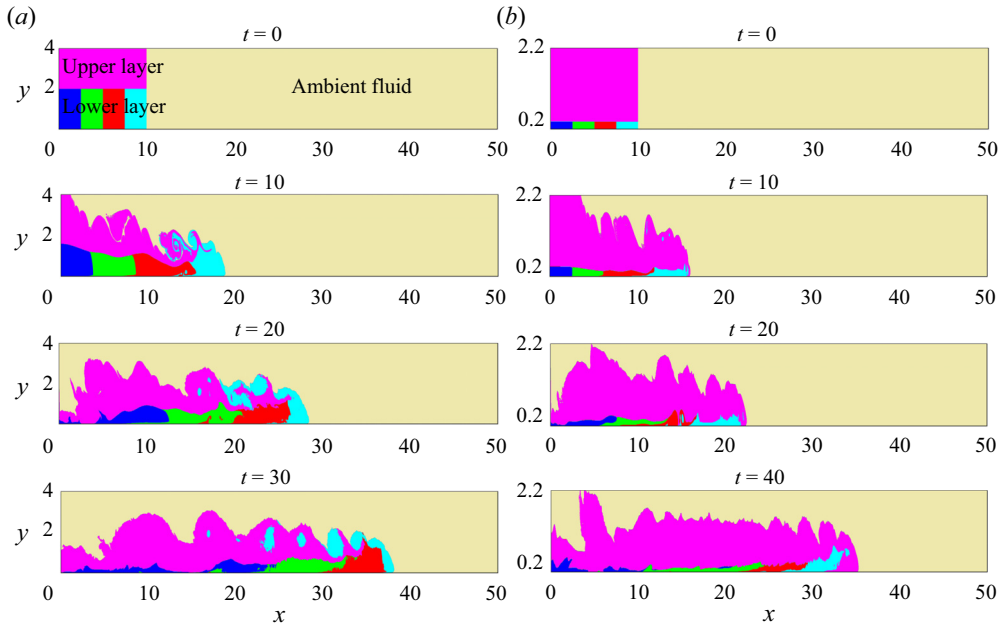


Figure 16. Temporal evolution of Lagrangian mixing between upper and lower layer fluids. (a) Run 9 ($R_\rho = 1.25$, $R_h = 1$): the lower layer fluid propagates faster and forms the head of the current. Significant mixing occurs in the wake of the head. (b) Run 2 ($R_\rho = 1.25$, $R_h = 0.1$): the upper layer fluid propagates faster and forms the current head. Relatively little mixing occurs between upper and lower layer fluids.

When the lower layer propagates faster, it rapidly forms a head after the gate is removed. The fluid in the head is carried upward by the ambient counterflow, so that it mixes with both the upper dense layer and the ambient fluid through the action of the interfacial Kelvin–Helmholtz vortices. The fluid within the head is constantly being replenished by the near-wall layer of the current, which moves faster than the head itself. It also entrains upper-layer fluid in a wedge-like shape behind the head, so that the entire head region becomes vigorously mixed, as shown in [figure 16\(a\)](#).

When the upper dense layer advances more rapidly than the lower layer, it surges ahead of the lower layer and forms the head of the current, cf. [figure 16\(b\)](#). Very little lower layer fluid enters the head under these conditions, and the two layers mix much less than in the flow of [figure 16\(a\)](#).

Previous studies have shown that the interfacial instabilities in stratified flow can help to predict the mixing properties between different layers (Balasubramanian & Zhong 2018; Martin, Negretti & Hopfinger 2019). Holmboe (1962) predicted Kelvin–Helmholtz instability and Holmboe instability at small and large Richardson numbers, respectively. The bulk Richardson number $J = g'\delta/(\Delta U)^2$ is defined, whereas Kelvin–Helmholtz instabilities can only occur when $J < 0.07$. Here, g' is the reduced gravity, $\delta = \Delta U/(\partial U/\partial y)_{max}$ is the shear layer thickness and ΔU is the velocity difference between the two layers or the current and the ambient fluid. Hazel (1972) found that Holmboe instabilities exist when $\delta/\eta > 2$, where $\eta = \Delta\rho/(\partial\rho/\partial y)_{max}$ is the density layer thickness, and $\Delta\rho$ is the density difference. We have measured J and δ/η from several profiles in different cases, which are presented in [table 3](#).

[Table 3](#) shows that the bulk Richardson numbers at the interface of the upper or lower layer and ambient fluid J_{U0} or J_{L0} are less than 0.07 for all cases. Kelvin–Helmholtz

Run	R_h	R_ρ	J_{U0}	J_{L0}	J_{UL}	δ_{U0}/η_{U0}	δ_{L0}/η_{L0}	δ_{UL}/η_{UL}
9	1.00	1.25	0.065	0.036	0.628	1.09	1.34	5.19
10	1.00	2.00	0.066	0.029	0.317	1.24	1.26	2.96
11	1.00	4.00	0.064	0.056	0.147	1.38	1.20	1.65
6	0.25	2.00	0.063	—	1.130	1.23	—	3.11
13	1.50	2.00	0.059	0.041	0.212	1.32	1.41	3.93

Table 3. Parameters J and δ/η for different cases.

instability is observed, as shown in figures 6 and 7. Note that for run 6, the upper and lower layers have similar front velocity, so a clear interface of the lower layer and ambient fluid does not exist. However, for the interface of the upper and lower layers, the bulk Richardson number J_{UL} is always larger than 0.07. In most of the cases δ_{UL}/η_{UL} is higher than 2, so that Holmboe instability occurs. In run 11, the large density ratio decreases the shear layer thickness, resulting in a small value of δ_{UL}/η_{UL} , which suppresses the Holmboe instability.

In order to quantify the mixing of upper and lower layer fluids as a function of R_ρ and R_h , we follow a computational approach similar to that of Tokyay, Constantinescu & Meiburg (2011), and evaluate the normalised volume V_m within which the concentrations of the upper and lower layer fluids both exceed a threshold value c_t :

$$V_m = \frac{1}{L_s \times h_U \times W} \int_0^H \int_0^L \int_0^W \gamma_{UL} \, dx \, dy \, dz, \tag{3.32}$$

$$\gamma_{UL} = \begin{cases} 1 & c_U > c_t \text{ and } c_L > c_t, \\ 0 & \text{otherwise.} \end{cases} \tag{3.33}$$

Figure 17 shows the temporal evolution of V_m for different values of R_ρ , R_h and $c_t = 0.02$. It demonstrates that R_ρ has a non-monotonic influence on the time history of V_m , so that at least during some flow stages mixing is most pronounced for intermediate values of R_ρ . These complex dynamics result from the competing effects of density stratification: on the one hand, strong stratification tends to suppress mixing; on the other hand, it increases the velocity difference between the layers, and thereby enhances the interfacial shear, which tends to promote mixing. At the same time, a larger velocity difference between the layers results in a head region that contains very little upper layer fluid, which in turn reduces mixing.

Figure 17 also indicates that a larger height ratio R_h uniformly tends to promote mixing, as it allows the dense fluid layers to release more potential energy, and to generate a more vigorous flow.

Similar to our approach in § 3.3, we evaluate the mixed volume V_m at $t = 20$ for different cases to show the quantitative dependence on the density ratio R_ρ and height ratio R_h . Figure 18(a) presents V_m as a function of R_h for different R_ρ . Here V_m increases with R_h , but is significantly reduced for large values of R_ρ . Figure 18(b) shows that for $R_\rho^{-0.25} R_h^{0.5}$ as the scaling parameter, the raw data can be well fitted ($R^2 = 0.962$) by a straight line of the form

$$V_m = 1.3 R_\rho^{-0.25} R_h^{0.5}. \tag{3.34}$$

Two-layer stratified gravity currents

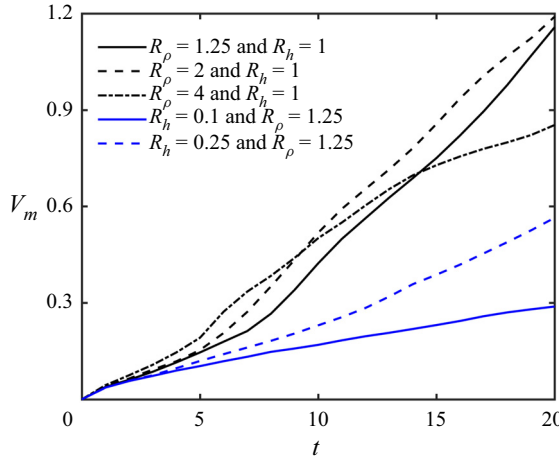


Figure 17. Time development of the mixed volume V_m for the height ratio $R_h = 1$ with different density ratios R_ρ (i.e. runs 9–11); and for the density ratio $R_\rho = 1.25$ with different height ratios R_h (i.e. runs 2, 5 and 9).

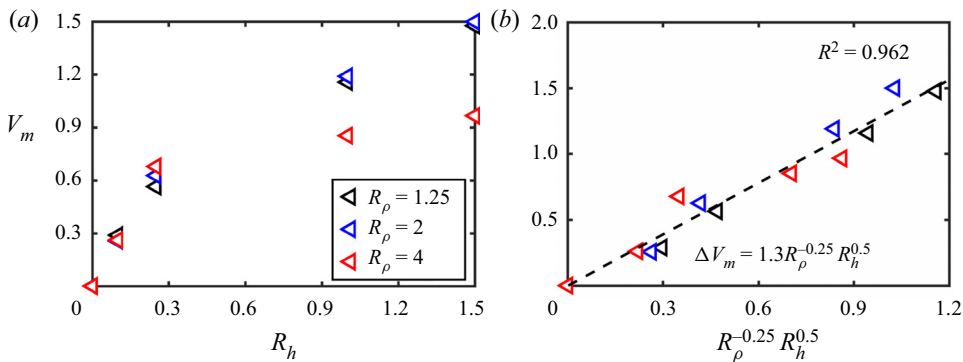


Figure 18. The mixed volume V_m for different R_ρ and R_h .

This equation is consistent with the fact that for $R_h = 0$ no mixed fluid is being generated. Although a larger R_h results in more mixing, an increase in R_ρ reduces the mixed volume V_m .

As the approach of Tokyay *et al.* (2011) is affected by the concentration threshold, we also present and compare the results of two alternate approaches for quantifying the mixing of the upper and lower layers. The first is based on the framework of Winters *et al.* (1995). In the closed system of our configuration, the increase in the background potential energy E_b results only from mixing, so that the evolution of E_b is a direct measure of the mixing. The background potential energy E_b is defined as

$$E_b = \int_0^W \int_0^H \int_0^L (c_U + R_L c_L)(y_* - h_L) dx dy dz, \tag{3.35}$$

where y_* is the vertical position in the state of the minimum potential energy. Note that here we also choose the initial top boundary of the lower dense layer $y = h_L$ as our reference level.

Figure 19 shows the time history of the normalised background potential energy $(E_b - E_{b,init})/E_{pU,init}$ for different cases. Here $E_{b,init}$ is the initial background potential energy.

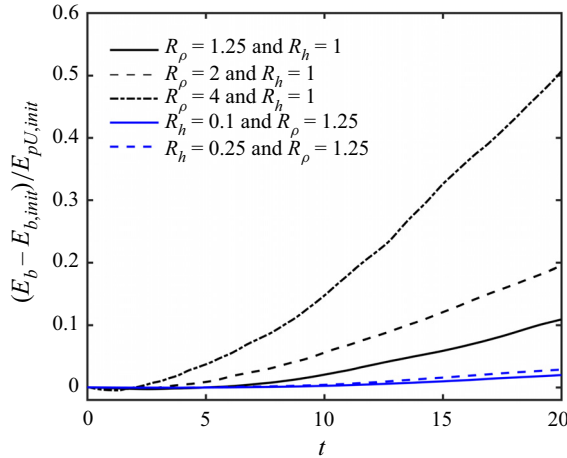


Figure 19. Time history of the normalised background potential energy $(E_b - E_{b,init})/E_{pU,init}$ for the height ratio $R_h = 1$ with different density ratios R_ρ (i.e. runs 9–11); and for the density ratio $R_\rho = 1.25$ with different height ratios R_h (i.e. runs 2, 5 and 9).

The energy E_b increases with time, which implies the presence of the irreversible mixing. However, different from the conclusion presented by figure 17, E_b monotonically increases with the density ratio R_ρ . This may be due to the fact that the density of the lower layer directly affects the calculation of the background potential energy. On the other hand, the evolution of E_b quantitatively presents the mixing not only between the upper and lower layers but also between the two layers and ambient fluid.

As a second approach, we first calculate the volumes that the upper layer, lower layer and the two-layer current occupy (García 1993). They are defined as V_U , V_L and V_{UL} :

$$V_U = \frac{\left(\int_0^W \int_0^H \int_0^L c_U \, dx \, dy \, dz \right)^2}{\int_0^W \int_0^H \int_0^L c_U^2 \, dx \, dy \, dz}, \quad (3.36)$$

$$V_L = \frac{\left(\int_0^W \int_0^H \int_0^L c_L \, dx \, dy \, dz \right)^2}{\int_0^W \int_0^H \int_0^L c_L^2 \, dx \, dy \, dz}, \quad (3.37)$$

$$V_{UL} = \frac{\left(\int_0^W \int_0^H \int_0^L c_U + c_L \, dx \, dy \, dz \right)^2}{\int_0^W \int_0^H \int_0^L (c_U + c_L)^2 \, dx \, dy \, dz}. \quad (3.38)$$

Then, the normalised mixed volume $V_{m,G}$ can be obtained as

$$V_{m,G} = \frac{V_U + V_L - V_{UL}}{L_s \times h_U \times W}. \quad (3.39)$$

Two-layer stratified gravity currents

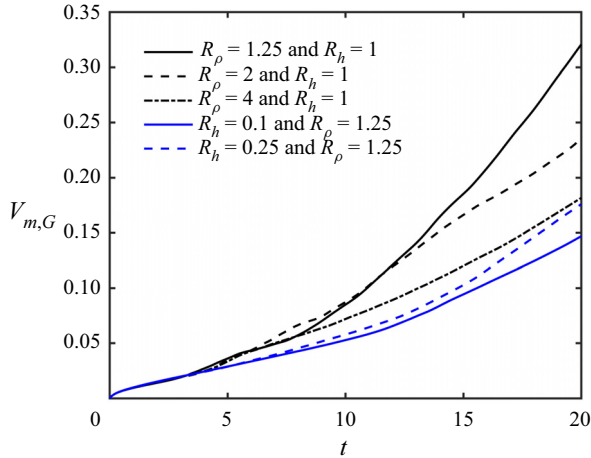


Figure 20. Time evolution of the normalised mixed volume $V_{m,G}$ for the height ratio $R_h = 1$ with different density ratios R_ρ (i.e. runs 9–11); and for the density ratio $R_\rho = 1.25$ with different height ratios R_h (i.e. runs 2, 5 and 9).

Figure 20 shows the time history of the normalised mixed volume $V_{m,G}$ for different cases. The values of $V_{m,G}$ are smaller than V_m shown in figure 17, which indicates that this method has a larger concentration threshold to define the boundary of the currents. At the same time, $V_{m,G}$ exhibits a monotonic decrease with the increase of the density ratio R_ρ . However, the values of V_m show the strongest mixing occurs for intermediate R_ρ , as larger values of R_ρ cause more low-concentration mixing.

3.5. Entrainment in two-layer gravity currents

As an essential feature, propagating gravity currents entrain ambient fluid, thus increasing their volume. The entrainment dynamics of two-layer currents are more complex than for single-layer currents, because they involve the interfaces of both layers with the ambient fluid. Following Ottolenghi *et al.* (2016), we define the entrainment velocity as the volume flux per unit area of the ambient fluid across the interface with the gravity currents:

$$U_e(t) = \frac{Q(t)}{0.5[S(t) + S(t - \Delta t)]}. \quad (3.40)$$

Here $S(t) = Wd_e(t)$ denotes the interfacial area at time t , where we define $d_e(t)$ as the rightmost location in the flow field with $c_U = 0.02$ or $c_L = 0.02$. Before the ambient counter-current reaches the left wall, $d_e(t)$ is the distance between this location and the front of the left-propagating ambient counter-current. Here $Q(t)$ is the instantaneous volume flux across the interface:

$$Q(t) = \frac{V(t) - V(t - \Delta t)}{\Delta t}, \quad (3.41)$$

where the volume of the two-layer gravity current is obtained as

$$V(t) = \int_0^W \int_0^H \int_0^L \gamma \, dx \, dy \, dz, \quad (3.42)$$

$$\gamma = \begin{cases} 1 & c_U > c_t \text{ or } c_L > c_t, \\ 0 & \text{otherwise.} \end{cases} \quad (3.43)$$

As a threshold value we take $c_t = 0.02$, consistent with Ottolenghi *et al.* (2016). The instantaneous entrainment parameter $E(t)$ of the current is then obtained as the ratio of the entrainment velocity to the front velocity of the two-layer gravity current relative to the ambient fluid:

$$E(t) = \frac{U_e(t)}{2U_f(t)}. \tag{3.44}$$

Here, we assume that the ambient fluid flows to the left with approximately $U_f(t)$, so that the velocity difference between the current and the ambient is $2U_f(t)$. The time-averaged entrainment behaviour of the flow is characterised by the bulk parameters:

$$\bar{Q}(t) = \frac{V(t) - V(0)}{t}, \tag{3.45}$$

$$\bar{U}_e(t) = \frac{\bar{Q}(t)}{0.5[S(t) + S(0)]}, \tag{3.46}$$

$$\bar{U}_f(t) = \frac{X_f(t) - X_f(0)}{t}, \tag{3.47}$$

$$\bar{E}(t) = \frac{\bar{U}_e(t)}{2\bar{U}_f(t)}. \tag{3.48}$$

Figure 21 compares the entrainment dynamics of single- (run 8 for $R_\rho = 1$ and $R_h = 1$) and two-layer (run 10 for $R_\rho = 2$ and $R_h = 1$) gravity currents. Figure 21(a,b) show the bulk and instantaneous entrainment fluxes, \bar{Q} and Q . Early on the dense fluid collapses and forms a well-defined head. Kelvin–Helmholtz billows evolve along the upper boundary and entrain ambient fluid, so that the bulk entrainment flux \bar{Q} increases with time. After the ambient counter-current reaches the left wall of the tank at $t \approx 12$, \bar{Q} levels off, with a larger value for the two-layer case. As compared with the bulk entrainment flux, the instantaneous flux Q exhibits more complex behaviour characterised by pronounced peaks. To establish the reasons for these peaks, we turn to the streamwise vorticity field averaged over the spanwise direction, cf. figure 22. Consistent with earlier work by Ottolenghi *et al.* (2016), we find that the Kelvin–Helmholtz billows trap the most ambient fluid when they reach their maximum size and intensity, resulting in the formation of instantaneous peaks in Q , cf. figure 22(a–d). The billows increase their size before the peak occurs, and they break down after that. Although the instantaneous fluxes Q evolve qualitatively similarly for the single- and two-layer cases, Q reaches a higher peak value for the two-layer case, cf. figure 22(e,f).

Figure 21(c,d) show the time evolution of the bulk and instantaneous entrainment velocities, \bar{U}_e and U_e . The high early values are related to the initial collapse and the formation of the head (He *et al.* 2018). After $t \approx 12$, \bar{U}_e and U_e decrease, because the tail region of the currents does not give rise to significant entrainment (Bhaganagar 2017). This is also reflected in the time evolution of the bulk and instantaneous entrainment parameters, \bar{E} and E , cf. figure 21(e,f). For both single- and two-layer flows we observe \bar{E} and E of $O(10^{-2})$, which is consistent with earlier investigations (Ottolenghi *et al.* 2016; He *et al.* 2018, 2019).

Figure 23 shows the influence of the density ratio R_ρ and the height ratio R_h on the evolution of the bulk entrainment parameter \bar{E} for two-layer gravity currents. Figure 23(a) demonstrates that \bar{E} decreases for larger R_ρ , since stronger stratification tends to dampen the Kelvin–Helmholtz instability while causing a larger front velocity. The effects of the

Two-layer stratified gravity currents

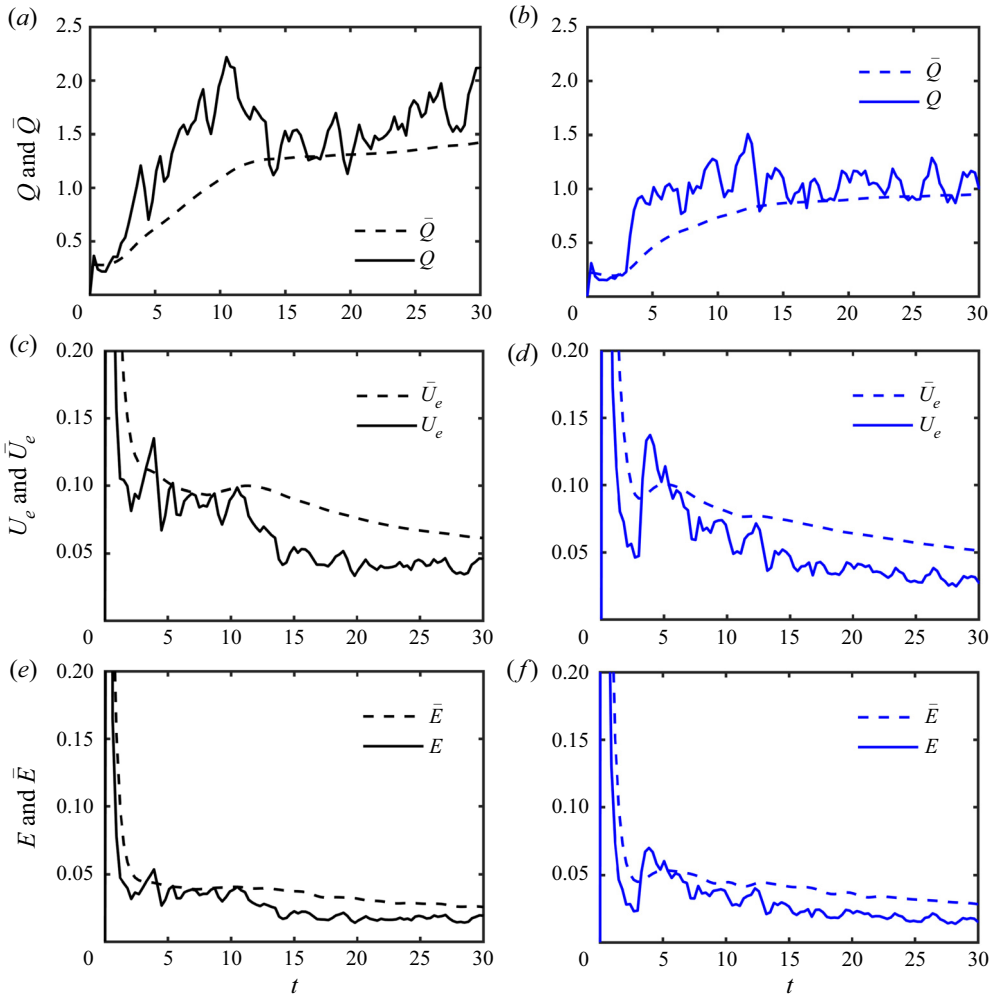


Figure 21. The time evolution of bulk and instantaneous entrainment properties for single- (run 8 for $R_\rho = 1$ and $R_h = 1$, blue lines, right column) and two-layer (run 10 for $R_\rho = 2$ and $R_h = 1$, black lines, left column) gravity currents: (a,b) bulk and instant entrainment fluxes \bar{Q} and Q ; (c,d) bulk and instantaneous entrainment velocities \bar{U}_e and U_e ; (e,f) bulk and instantaneous entrainment parameters \bar{E} and E .

layer height ratio R_h are analysed in [figure 23\(b\)](#). Until the left-propagating counter-current encounters the left wall, R_h has only a minor influence on \bar{E} , because the entrainment is dominated by the lower-layer current. After the reflection of the counter-current from the left wall, however, it is evident that larger R_h -values enhance entrainment, so that \bar{E} increases.

3.6. Constant excess weight

As mentioned in § 2.1, so far we have employed the thickness and density of the upper layer to scale the simulations. Consequently, the total excess weight in the initial condition has not been held constant for the different cases. In this section, we keep the total excess

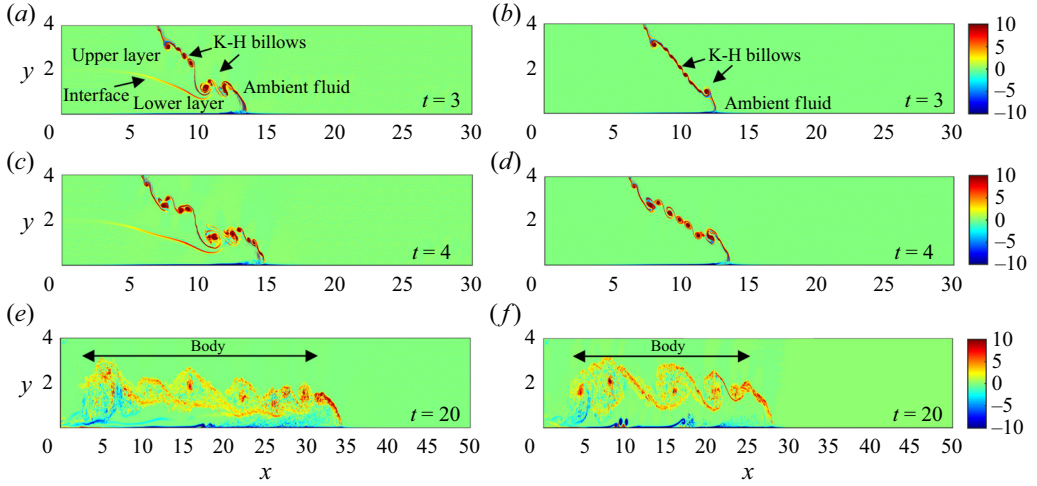


Figure 22. The dimensionless vorticity fields ω_{xy} of single- (run 8 for $R_\rho = 1$ and $R_h = 1$) and two-layer (run 10 for $R_\rho = 2$ and $R_h = 1$) gravity currents: (a,c,e) for the two-layer case; (b,d,f) for the single-layer case.

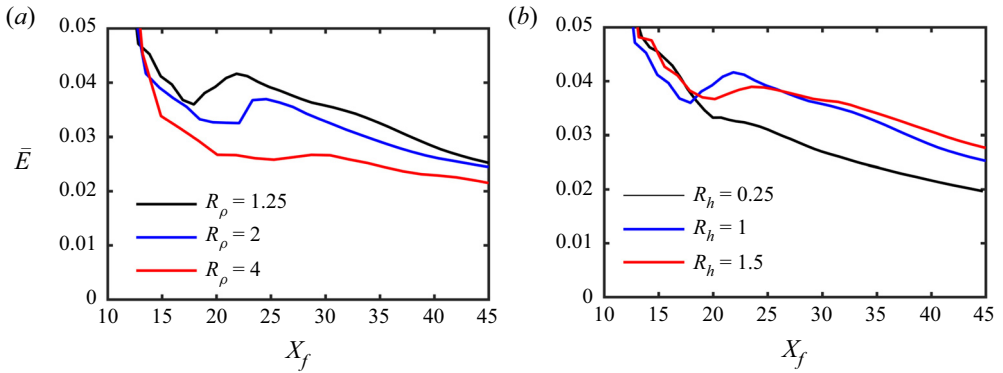


Figure 23. The evolution of bulk entrainment parameter \bar{E} for two-layer gravity currents as a function of front location X_f with different R_ρ and R_h : (a) $R_h = 1$ with different R_ρ , i.e. runs 9–11; (b) $R_\rho = 1.25$ with different R_h , i.e. runs 5, 9 and 12.

weight constant and run several cases with different density and height ratios of the upper and lower layers.

We introduce characteristic length and velocity scales $\hat{H}/2$ and $\hat{u}_{b,A} = \sqrt{\hat{g}\hat{H}(\hat{\rho}_A - \hat{\rho}_0)/2\hat{\rho}_0}$ to non-dimensionalise the governing equations. Here $\hat{\rho}_A = (\hat{\rho}_U\hat{h}_U + \hat{\rho}_L\hat{h}_L)/\hat{H}$ is the average density of the lock fluid. The non-dimensional (2.10)–(2.12) can be rewritten as

$$\frac{\partial u_i}{\partial t} + u_j \frac{\partial u_i}{\partial x_j} = -\frac{\partial p}{\partial x_i} + \frac{1}{Re_A} \frac{\partial^2 u_i}{\partial x_j \partial x_j} + (R_{UCU} + R_{LC L}) \mathbf{e}_i^g, \quad (3.49)$$

Run	h_U	h_L	R_h	R_U	R_L	R_ρ	Re_A
15	1.0	1.0	1.00	1.00	1.00	1.0	7000
16	1.0	1.0	1.00	0.80	1.20	1.5	7000
17	1.0	1.0	1.00	0.50	1.50	3.0	7000
18	1.6	0.4	0.25	0.91	1.36	1.5	7000
19	0.4	1.6	4.00	0.71	1.07	1.5	7000

Table 4. Simulation cases and parameters.

$$\frac{\partial c_U}{\partial t} + u_j \frac{\partial c_U}{\partial x_j} = \frac{1}{Pe_A} \frac{\partial^2 c_U}{\partial x_j \partial x_j}, \tag{3.50}$$

$$\frac{\partial c_L}{\partial t} + u_j \frac{\partial c_L}{\partial x_j} = \frac{1}{Pe_A} \frac{\partial^2 c_L}{\partial x_j \partial x_j}. \tag{3.51}$$

The dimensionless parameters are the Reynolds number Re_A ,

$$Re_A = \frac{\hat{u}_{b,A} \hat{H}}{2\hat{\nu}}, \tag{3.52}$$

the Péclet number Pe_A ,

$$Pe_A = \frac{\hat{u}_{b,A} \hat{H}}{2\hat{\kappa}}, \tag{3.53}$$

and the density parameters for the upper layer R_U and lower layer R_L ,

$$R_U = \frac{\hat{\rho}_U - \hat{\rho}_0}{\hat{\rho}_A - \hat{\rho}_0}, \tag{3.54}$$

$$R_L = \frac{\hat{\rho}_L - \hat{\rho}_0}{\hat{\rho}_A - \hat{\rho}_0}. \tag{3.55}$$

The initial and boundary conditions are as described in § 2.2. The parameters of the simulations are listed in table 4. All cases have the same averaged density of the two-layer lock fluids.

Figure 24 shows the evolution of the two-layer gravity currents with different density ratio R_ρ and height ratio R_h . Although having different stratification, the two-layer currents propagate at comparable velocities, due to the identical initial averaged density ρ_A . Figure 24(a–c) compare three cases with constant height ratio $R_h = 1$ but different density ratio R_ρ , i.e. runs 15–17. With increasing R_ρ , the upper layer slows down and the main body of the currents becomes thinner. Figure 24(d–f) present the flows with identical density ratio $R_\rho = 1.5$ and different height ratio, i.e. runs 16, 18 and 19. When increasing R_h , the lower layer has a larger volume but smaller density. For a small height ratio $R_h = 0.25$, the fronts of the upper and lower layers have similar velocities. For larger height ratios ($R_h = 1$ and 4), the lower layer propagates faster than the upper layer.

4. Conclusions

We have employed three-dimensional direct numerical simulations to investigate the mixing and entrainment dynamics, as well as the energetics, of gravity currents emerging from a two-layer stratified lock. The dominant governing dimensionless parameters take

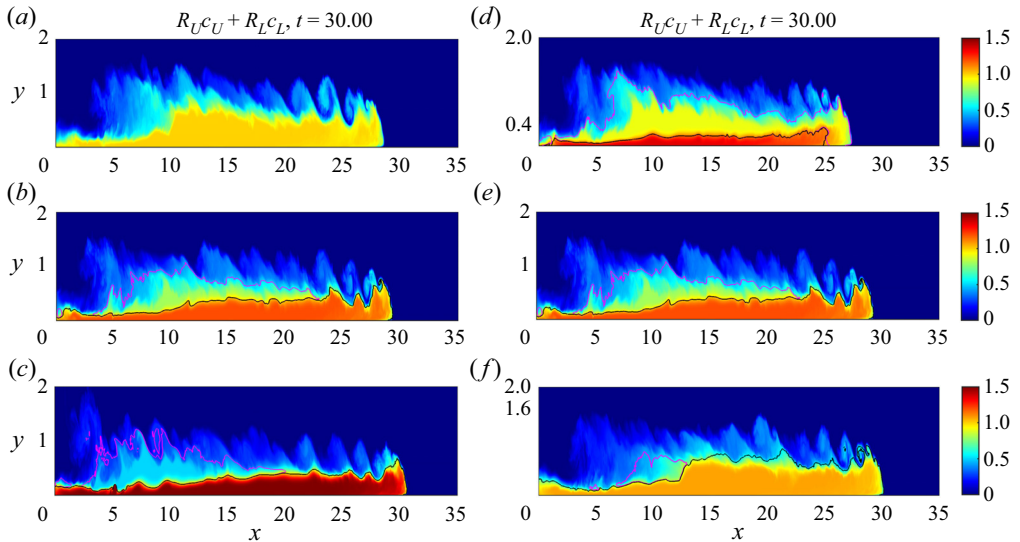


Figure 24. Instantaneous concentration field, along with the concentration isosurfaces $c_U = 0.5$ and $c_L = 0.5$ for both upper and lower layers at $t = 30$: (a–c) runs 15–17 for $R_h = 1$ and $R_\rho = 1, 1.5$ and 3; (d–f) runs 18, 16 and 19 for $R_\rho = 1.5$ and $R_h = 0.25, 1$ and 4.

the form of a layer thickness ratio R_h and a density ratio R_ρ . Depending on the values of these parameters, the simulations show that either the upper layer or lower layer fluid can propagate faster, so that it forms the head of the overall gravity current. Similarly, the simulations clarify the role of R_h and R_ρ in determining the long-term density structure of the gravity current, which can range from strongly stratified to near-complete mixing. These simulation results outline possible scenarios for the behaviour of certain classes of turbidity currents and powder snow avalanches, with a dense, concentrated layer near the base and a more dilute layer above. In particular, the simulations demonstrate that depending on the specific flow conditions, either the dense or the dilute layer can propagate more rapidly.

Building on the earlier circulation-based approach introduced by Borden & Meiburg (2013), we propose a theoretical model for predicting the overall gravity current height, along with the front velocity of the two layers, when the lower layer moves faster than the upper layer. This model clarifies the role of the height and thickness ratios in determining the velocity structure of the current, and it sheds light on the dynamics of the ambient counter-current as it slows down the propagation of the upper gravity current layer.

By conducting a detailed analysis of the energy budget, we have been able to quantify the conversion between potential and kinetic energy as a function of the parameters R_ρ and R_h . This analysis also clarifies the energy transfer rates between the different layers of the gravity current, and from the gravity current to the ambient fluid. In particular, it shows that the dense gravity current layer may gain or lose energy, depending on the specific flow conditions.

Further work is needed in order to clarify how the present results may change in the presence of a sloping bottom, or how they may be modified if the scalar field has a settling velocity associated with it, if it exchanges mass with the bed through erosion or deposition, or if it gives rise to non-Newtonian rheology.

Funding. Funding for this work was provided under NSF grant CBET-2138583, under the National Natural Science Foundation of China (52171276), as well as by the Key Research and Development Program of Zhejiang Province (2021C03180). Computational resources for this work used the Extreme Science and Engineering Discovery Environment (XSEDE), supported by the National Science Foundation, USA, grant no. TG-CTS150053. R.Z. thanks the China Scholarship Council for providing him with a scholarship to study at UCSB.

Declaration of interests. The authors report no conflict of interest.

Author ORCIDs.

 Zhiguo He <https://orcid.org/0000-0002-0612-9062>;

 Eckart Meiburg <https://orcid.org/0000-0003-3670-8193>.

REFERENCES

- ADDUCE, C., SCIORTINO, G. & PROIETTI, S. 2012 Gravity currents produced by lock exchanges: experiments and simulations with a two-layer shallow-water model with entrainment. *J. Hydraul. Engng* **138** (2), 111–121.
- BALASUBRAMANIAN, S. & ZHONG, Q. 2018 Entrainment and mixing in lock-exchange gravity currents using simultaneous velocity-density measurements. *Phys. Fluids* **30** (5), 056601.
- BHAGANAGAR, K. 2017 Role of head of turbulent 3-d density currents in mixing during slumping regime. *Phys. Fluids* **29** (2), 020703.
- BIEGERT, E., VOWINCKEL, B. & MEIBURG, E. 2017a A collision model for grain-resolving simulations of flows over dense, mobile, polydisperse granular sediment beds. *J. Comput. Phys.* **340**, 105–127.
- BIEGERT, E., VOWINCKEL, B., OUILLOIN, R. & MEIBURG, E. 2017b High-resolution simulations of turbidity currents. *Prog. Earth Planet. Sci.* **4** (1), 33.
- BIRMAN, V.K., MARTIN, J.E. & MEIBURG, E. 2005 The non-Boussinesq lock-exchange problem. Part 2. High-resolution simulations. *J. Fluid Mech.* **537**, 125–144.
- BORDEN, Z. & MEIBURG, E. 2013 Circulation based models for Boussinesq gravity currents. *Phys. Fluids* **25** (10), 101301.
- BUCKEE, C., KNELLER, B. & PEAKALL, J. 2001 Turbulence structure in steady, solute-driven gravity currents. In *Particulate Gravity Currents* (ed. W. McCaffrey, B. Kneller & J. Peakall), pp. 173–187. Wiley.
- CANTERO, M.I., BALACHANDAR, S., CANTELLI, A., PIRMEZ, C. & PARKER, G. 2009 Turbidity current with a roof: direct numerical simulation of self-stratified turbulent channel flow driven by suspended sediment. *J. Geophys. Res.* **114**, C03008.
- DAI, A. 2017 Experiments on two-layer density-stratified inertial gravity currents. *Phys. Rev. Fluids* **2** (7), 073802.
- ESPATH, L.F.R., PINTO, L.C., LAIZET, S. & SILVESTRINI, J.H. 2014 Two- and three-dimensional direct numerical simulation of particle-laden gravity currents. *Comput. Geosci.* **63**, 9–16.
- FRAGOSO, A.T., PATTERSON, M.D. & WETTLAUFER, J.S. 2013 Mixing in gravity currents. *J. Fluid Mech.* **734**, R2.
- GARCÍA, M.H. 1993 Hydraulic jumps in sediment-driven bottom currents. *J. Hydraul. Engng* **119** (10), 1094–1117.
- GERBER, G., DIEDERICKS, G. & BASSON, G.R. 2011 Particle image velocimetry measurements and numerical modeling of a saline density current. *J. Hydraul. Engng* **137** (3), 333–342.
- GLADSTONE, C., RITCHIE, L.J., SPARKS, R.S.J. & WOODS, A.W. 2004 An experimental investigation of density-stratified inertial gravity currents. *Sedimentology* **51** (4), 767–789.
- HACKER, J., LINDEN, P.F. & DALZIEL, S.B. 1996 Mixing in lock-release gravity currents. *Dyn. Atmos. Oceans* **24** (1–4), 183–195.
- HALLWORTH, M.A., HUPPERT, H.E., PHILLIPS, J.C. & SPARKS, R.S.J. 1996 Entrainment into two-dimensional and axisymmetric turbulent gravity currents. *J. Fluid Mech.* **308**, 289–311.
- HAMPTON, M.A. 1972 The role of subaqueous debris flow in generating turbidity currents. *J. Sedim. Res.* **42** (4), 775–793.
- HÄRTEL, C., KLEISER, L., MICHAUD, M. & STEIN, C.F. 1997 A direct numerical simulation approach to the study of intrusion fronts. *J. Engng Maths* **32** (2), 103–120.
- HÄRTEL, C., MEIBURG, E. & NECKER, F. 2000 Analysis and direct numerical simulation of the flow at a gravity-current head. Part 1. Flow topology and front speed for slip and no-slip boundaries. *J. Fluid Mech.* **418**, 189–212.

- HAZEL, P. 1972 Numerical studies of the stability of inviscid stratified shear flows. *J. Fluid Mech.* **51** (1), 39–61.
- HE, Z., ZHAO, L., HU, P., YU, C.-H. & LIN, Y.-T. 2018 Investigations of dynamic behaviors of lock-exchange turbidity currents down a slope based on direct numerical simulation. *Adv. Water Resour.* **119**, 164–177.
- HE, Z., ZHAO, L., LIN, T., HU, P., LV, Y., HO, H.-C. & LIN, Y.-T. 2017 Hydrodynamics of gravity currents down a ramp in linearly stratified environments. *J. Hydraul. Engng* **143** (3), 04016085.
- HE, Z., ZHAO, L., ZHU, R. & HU, P. 2019 Separation of particle-laden gravity currents down a slope in linearly stratified environments. *Phys. Fluids* **31**, 106602.
- HE, Z., ZHU, R., ZHAO, L., CHEN, J., LIN, Y.-T. & YUAN, Y. 2021 Hydrodynamics of weakly and strongly stratified two-layer lock–release gravity currents. *J. Hydraul. Res.* **59** (6), 989–1003.
- HOLMBOE, J. 1962 On the behavior of symmetric waves in stratified shear layers. *Geophys. Publ.* **24**, 67–113.
- KNELLER, B., NASR-AZADANI, M.M., RADHAKRISHNAN, S. & MEIBURG, E. 2016 Long-range sediment transport in the world’s oceans by stably stratified turbidity currents. *J. Geophys. Res.* **121** (12), 8608–8620.
- MARTIN, A., NEGRETTI, M.E. & HOPFINGER, E.J. 2019 Development of gravity currents on slopes under different interfacial instability conditions. *J. Fluid Mech.* **880**, 180–208.
- MEIBURG, E. & KNELLER, B. 2010 Turbidity currents and their deposits. *Annu. Rev. Fluid Mech.* **42**, 135–156.
- NASR-AZADANI, M.M., HALL, B. & MEIBURG, E. 2013 Polydisperse turbidity currents propagating over complex topography: comparison of experimental and depth-resolved simulation results. *Comput. Geosci.* **53**, 141–153.
- NASR-AZADANI, M.M. & MEIBURG, E. 2011 TURBINS: an immersed boundary, Navier–Stokes code for the simulation of gravity and turbidity currents interacting with complex topographies. *Comput. Fluids* **45** (1), 14–28.
- NECKER, F., HÄRTEL, C., KLEISER, L. & MEIBURG, E. 2002 High-resolution simulations of particle-driven gravity currents. *Intl J. Multiphase Flow* **28** (2), 279–300.
- NECKER, F., HÄRTEL, C., KLEISER, L. & MEIBURG, E. 2005 Mixing and dissipation in particle-driven gravity currents. *J. Fluid Mech.* **545**, 339–372.
- NOGUEIRA, H.I.S., ADDUCE, C., ALVES, E. & FRANCA, M.J. 2014 Dynamics of the head of gravity currents. *Environ. Fluid Mech.* **14** (2), 519–540.
- OOI, S.K., CONSTANTINESCU, G. & WEBER, L. 2009 Numerical simulations of lock-exchange compositional gravity current. *J. Fluid Mech.* **635**, 361–388.
- OTTOLENGHI, L., ADDUCE, C., INGHILESI, R., ARMENIO, V. & ROMAN, F. 2016 Entrainment and mixing in unsteady gravity currents. *J. Hydraul. Res.* **54** (5), 541–557.
- ROSS, A.N., DALZIEL, S.B. & LINDEN, P.F. 2006 Axisymmetric gravity currents on a cone. *J. Fluid Mech.* **565**, 227–253.
- ROTTMAN, J.W. & SIMPSON, J.E. 1983 Gravity currents produced by instantaneous releases of a heavy fluid in a rectangular channel. *J. Fluid Mech.* **135**, 95–110.
- SAMASIRI, P. & WOODS, A.W. 2015 Mixing in axisymmetric gravity currents. *J. Fluid Mech.* **782**, R1.
- SEQUEIROS, O.E., SPINOWINE, B., BEAUBOUF, R.T., SUN, T., GARCÍA, M.H. & PARKER, G. 2010 Characteristics of velocity and excess density profiles of saline underflows and turbidity currents flowing over a mobile bed. *J. Hydraul. Engng* **136** (7), 412–433.
- SHER, D. & WOODS, A.W. 2015 Gravity currents: entrainment, stratification and self-similarity. *J. Fluid Mech.* **784**, 130–162.
- SIMPSON, J.E. 1982 Gravity currents in the laboratory, atmosphere, and ocean. *Annu. Rev. Fluid Mech.* **14** (1), 213–234.
- TOKYAY, T., CONSTANTINESCU, G. & MEIBURG, E. 2011 Lock-exchange gravity currents with a high volume of release propagating over a periodic array of obstacles. *J. Fluid Mech.* **672**, 570–605.
- WINTERS, K.B., LOMBARD, P.N., RILEY, J.J. & D’ASARO, E.A. 1995 Available potential energy and mixing in density-stratified fluids. *J. Fluid Mech.* **289**, 115–128.
- WU, C.-S. & DAI, A. 2020 Experiments on two-layer stratified gravity currents in the slumping phase. *J. Hydraul. Res.* **58** (5), 831–844.
- ZEMACH, T. & UNGARISH, M. 2020 A model for the propagation of inertial gravity currents released from a two-layer stratified lock. *J. Fluid Mech.* **903**, R5.
- ZEMACH, T. & UNGARISH, M. 2021 Gravity currents with internal stratification in channels of non-rectangular cross-section. *Eur. J. Mech. B/Fluids* **89**, 83–92.
- ZHAO, L., HE, Z., LV, Y., LIN, Y.-T., HU, P. & PÄHTZ, T. 2018 Front velocity and front location of lock-exchange gravity currents descending a slope in a linearly stratified environment. *J. Hydraul. Engng* **144** (11), 04018068.
- ZHU, R., HE, Z. & MEIBURG, E. 2021 Removal of a dense bottom layer by a gravity current. *J. Fluid Mech.* **916**, A30.