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# Eocene giant ants, Arctic intercontinental dispersal, and hyperthermals revisited: discovery of fossil *Titanomyrma* (Hymenoptera: Formicidae: Formiciinae) in the cool uplands of British Columbia, Canada

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## Abstract

We examine the implications for intercontinental dispersal of the extinct ant genus, *Titanomyrma* Archibald *et al.* (Hymenoptera: Formicidae: Formiciinae), following the discovery of its first fossil in Eocene temperate upland Canada. Modern Holarctic distributions of plants and animals were in part formed by dispersals across Late Cretaceous through early Eocene Arctic land bridges. Mild winters in a microthermal Arctic would allow taxa today restricted to the tropics by cold intolerance to cross, with episodic hyperthermal events allowing tropical taxa requiring hot climates to cross. Modern ants with the largest queens inhabit low latitudes of high temperature and mild coldest months, whereas those with smaller queens inhabit a wide variety of latitudes and climates. Gigantic and smaller formiciine ants (*Titanomyrma* and *Formicium* Westwood) are known from Europe and North America in the Eocene. The new Canadian *Titanomyrma* inhabited a cooler upland. It is incomplete, indistinctly preserved, and distorted in fossilisation, and so we do not assign it to a species or erect a new one for it. The true size of this fossil is unclear by this distortion: small size would support gigantism in *Titanomyrma* requiring hot climates and dispersal during hyperthermals; if it was large, it may have been cold-winter intolerant and able to have crossed during any time when the land bridge was present.

## Introduction

Modern Holarctic distributions of plants and animals were formed in part by the existence of intercontinental dispersal corridors in the Late Cretaceous through early Paleogene, together with a change in climate relative to the tolerances of organisms that might expand their ranges across them. Land connections between North America and Europe *via* Greenland were present in the Maastrichtian through late Danian by the northern de Geer route to Fennoscandia, and in the Thanetian and early Ypresian by a southern Thulean route through Greenland, the Faroes, and the United Kingdom (Brikiatis 2014). North America and East Asia were episodically

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connected across Beringia in the Danian, Thanetian, and possibly the Eocene. This was a time of “Greenhouse World” climate, with high global mean annual temperatures, low pole-to-equator mean annual temperature gradients, poles free of ice sheets, and low extra-tropical temperature seasonality resulting in mild winters in cooler higher latitudes and elevations (Zachos *et al.* 2008, 2010; Archibald *et al.* 2010; National Research Council 2011).

In contrast, our post-Eocene “Icehouse World” climatic regime is characterised by cooler mean annual temperatures, a steeper equator-to-pole mean annual temperature gradient, and greater temperature seasonality outside of the tropics (Zachos *et al.* 2008, 2010; Archibald *et al.* 2010; National Research Council 2011). Thermophilic organisms have become restricted to lower latitudes because they require the high mean annual temperature values found there. Others now restricted to lower latitudes are cryophobic: they do not require high mean annual temperatures but are intolerant of cold extra-tropical winters associated with increased temperature seasonality (Archibald and Farrell 2003; Archibald *et al.* 2010). In the Greenhouse World, however, these two groups partially separated, with cryophobic organisms ranging into cooler higher latitudes and elevations with mild winters – that is, regions with lowered mean annual temperatures that would exclude thermophilic organisms. This allowed the existence of communities without modern analogues – for example, forests with palms, alders, maples, and spruce in upper microthermal localities of the Ypresian Okanagan Highlands. These high-paleoelevation localities in far-western mid-latitude North America (*e.g.*, Archibald and Farrell 2003; Archibald *et al.* 2010, 2014) reveal clear examples of such mixed climatic-preference plant and insect communities.

This climatic regime and community response has implications for Greenhouse World trans-Arctic dispersal. Analysis by the TEX<sub>86</sub> proxy thermometer method gives megathermal background sea surface mean annual temperature values for the Ypresian Arctic, but this method may produce misleadingly high temperature values (Kim *et al.* 2010; Sluijs *et al.* 2020). Lower estimates of upper microthermal terrestrial mean annual temperatures are given for the Ypresian Arctic using paleobotanical proxies, like the mean annual temperatures determined for most localities of the montane Okanagan Highlands by the same methods (Greenwood *et al.* 2005; Sunderlin *et al.* 2011; Mathewes *et al.* 2016; West *et al.* 2020). In the Eocene Arctic, as in the Okanagan Highlands, a suite of frost-intolerant plants and animals indicates mild winters with few, if any, frost days (*e.g.*, Wing and Greenwood 1993; Archibald and Farrell 2003; Archibald *et al.* 2010, 2011a, 2014; Eberle and Greenwood 2012). A microthermal Arctic intercontinental dispersal corridor with mild winters would then have acted as a filter, allowing passage of cryophobic taxa while excluding thermophilic taxa.

Throughout the Ypresian (47.8–56.0 Ma), this temperate and equable Arctic climate was, however, punctuated by brief hyperthermal events of global warming associated with large quantities of carbon injected into the atmosphere–ocean system. The most notable hyperthermal was the Paleocene–Eocene Thermal Maximum, with a release of more than 2000 gigatonnes of carbon commencing at the Paleocene–Eocene boundary and lasting less than 20 ky, decreasing to background values in about 200 ky, with a peak increase in global temperature of 5–8 °C (Wing *et al.* 2005; Zachos *et al.* 2008; McInerney and Wing 2011; Westerhold *et al.* 2017; Stokke *et al.* 2020; Vickers *et al.* 2020; Reinhardt *et al.* 2022). Such hyperthermals would have acted as climatic “gates” for thermophilic organisms, briefly opening periodically to allow high-latitude passage between continents.

Archibald *et al.* (2011b) examined this “bridges with gates” dispersal model regarding size and habitat climate in ants, comparing modern species that have queens of at least 3 cm in length with those of the giant Eocene ants of the extinct subfamily Formiciinae (Hymenoptera: Formicidae) (species of *Titanomyrma* Archibald *et al.* and the parataxon *Formicium* Westwood, those known only from wings). They found that modern ants with the largest queens inhabit only low latitudes,

almost exclusively in the tropics, and so great size of queens is associated with being either thermophilic or cryophobic.

The queens of giant species of Formiciine are by far the largest of any extant or fossil Hymenoptera: *T. gigantea* (Lutz) and *T. simillima* (Lutz) (Ypresian Messel and Lutetian Eckfeld Maar, Germany), *T. lubei* Archibald *et al.* (Ypresian Green River Formation, Wyoming, United States of America), and *F. mirabile* (Cockerell) (Lutetian Bournemouth Group, United Kingdom). *Titanomyrma gigantea* queens reached 7 cm in length and possessed a wingspan of about 16 cm (Lutz 1990). Archibald *et al.* (2011b) found that these are known only at localities of high mesothermal to megathermal mean annual temperatures and never at localities of cooler mean annual temperatures such as the Okanagan Highlands, which is consistent with their being thermophilic, not cryophobic. Smaller Formiciinae are also known from megathermal localities (*F. brodiei* Westwood, Ypresian–Lutetian Bournemouth Group, United Kingdom) or of undetermined mean annual temperatures (*F. berryi* Carpenter, middle Eocene Claiborne Formation, Tennessee, United States of America; Archibald *et al.* 2011b). Modern ant species with smaller queens are known from all latitudes where ants are found. The discovery of giant formiciine ants in the Okanagan Highlands would imply that they were cryophobic, not thermophilic. Smaller formiciine there would be consistent with species with giant queens being thermophilic and constrained to higher mean annual temperature regions, and those with smaller queens being distributed across a variety of climates, including regions with lower mean annual temperatures.

Both giant and smaller Formiciinae appear in the second half of the Ypresian in Europe and North America, implying a cross-Arctic dispersal of both size groups in either direction at least once sometime in the Late Cretaceous through early Ypresian under Greenhouse World conditions with episodic hyperthermals. Dispersal of one size group with separate origins of the other on both sides of the Atlantic seems less likely.

Here, we examine the implications for trans-Arctic dispersal of a new *Titanomyrma* from the microthermal Okanagan Highlands Allenby Formation of British Columbia, Canada.

## Material and methods

We examined a fossil from the Allenby Formation near Princeton, British Columbia, Canada, in the collections of the Beaty Biodiversity Museum (Vancouver, British Columbia, Canada) and compared it with other Formiciinae, including two new specimens from the Green River Formation of Wyoming, United States of America in the collections of Fossil Butte National Monument (Kemmerer, Wyoming, United States of America).

Photography of the Allenby ant was done at the Parks Canada laboratory, Vancouver, British Columbia, Canada, using a Zeiss microphotography system (Oberkochen, Germany) and of the Green River specimens at the University of Colorado, Boulder, Colorado, United States of America and the Denver Museum of Nature & Science, Denver, Colorado.

We follow the morphological terminology of Lutz (1986, 1990) and Wappler (2003), except that we refer to abdominal, not gaster segment numbers (*e.g.*, abdominal segments A3–7 are gaster segments I–V). Contrary character states in compared taxa are provided in brackets.

We refer to the general mean annual temperature categories of Wolfe (1975), rather than estimated values that vary by a few degrees in differing analyses (various equations for leaf margin analyses, the Climate Leaf Analysis Multivariate Program, and nearest living relatives of plants analysis; *e.g.*, Greenwood *et al.* 2005), as this level of precision is relevant here. These mean annual temperature categories are as follows: microthermal:  $\leq 13$  °C; mesothermal:  $> 13$  °C,  $< 20$  °C; and megathermal:  $\geq 20$  °C.

## Localities

**The Allenby Formation.** This formation is part of the Okanagan Highland series of fossiliferous lacustrine basins scattered across about 1000 km from west-central British Columbia into northern Washington, United States of America (Read 2000; Archibald *et al.* 2011a). The ant was collected in an exposure of Vermilion Bluffs Shale Unit about 4 km southwest of the village of Princeton, British Columbia. A nearby Vermilion Bluffs Unit outcrop (the Billy's Family Restaurant locality) is estimated by U–Pb decay to be Ypresian,  $51.85 \pm 0.85$  Ma (Rubino *et al.* 2021). The Allenby Formation forest had an estimated upper microthermal mean annual temperature (Greenwood *et al.* 2005; West *et al.* 2020), with few or no frost days in its coldest month, as indicated by a suite of frost-intolerant plants such as the coryphoid palm *Uhlia allenbyensis* Erwin and Stockey (Arecaceae) (Erwin and Stockey 1991; Archibald *et al.* 2014).

**The Green River Formation.** This formation includes lacustrine sediments deposited in Wyoming, Colorado, and Utah, United States of America (Smith and Carroll 2015) in three Ypresian lakes. The ants were found in the laminated limestones of the White Marker Unit near the base of the Angelo Member of the Fossil Lake deposits in Fossil Basin, Wyoming (Buchheim *et al.* 2011). This unit, approximately 15 cm thick, lies about 5.5 m above the K-spar tuff, which has a radiometric age of  $51.98 \pm 0.35$  Ma (Smith and Carroll 2015). Green River shales were deposited in a lowland intermontane basin with high mesothermal to mostly megathermal mean annual temperature values, as determined by paleobotanical and isotopic analyses (Wilf 2000; Archibald *et al.* 2011b).

## Specimens

The Allenby ant, new Green River Formation ants, and the only specimen of *T. lubei* vary in preservation and cannot be confidently considered separate species (unless the Allenby ant is small; see below). We, therefore, treat the new fossils as *Titanomyrma* sp. pending more complete and undistorted specimens. The Allenby ant is a formicine by the “crowding” of the forewing pterostigma and cells 1–2R, 1M, and 1Rs – that is, these are small and grouped in the anterior middle wing – and this Allenby specimen and the new Green River specimens are *Titanomyrma* species by their diagnostic slit-like, prominent spiracles (Lutz 1986).

**The Allenby Formation ant.** Beaty Biodiversity Museum collection BBM-PAL-2022-00001 (part only), queen

(Figs. 1, 2A, 3, 4A).

Measurements are not given by uncertainty due to distortion; shapes are described after graphic adjustment (see Fig. 2 and size discussion, below). Head: with rounded posterior margins, not lobed, missing antennae, capsule about 2/5 alitrunk length; eyes oval, maximum length oriented anterodorsally, located at cephalic mid-length; large, triangular mandibles, about half head length, 8–10 coarse teeth. Alitrunk: poorly preserved. Forewing (basal and distal-posterior portions not known), with venation of preserved portions most like that of *T. simillima* but differs by: 3-Rs joins M approximately middle between 1-M, m-cu [*T. simillima*: joins close to m-cu], cell 1Rs evenly widens distally (to rs-m) [narrows distad 2r], cell 1M about as high (Cu to Rs+M) as wide (Cu) [about half as high as wide], cu-a joins M+Cu at approximately half its length [joins very close to 1-M, Cu]. Hind wings missing. Legs: portions preserved short, stout. Waist: single segmented, petiole lacking anterior peduncle, otherwise indistinctly preserved. Partial propodeal spiracle preserved, slit-like, long. Gaster: only the first segment (A3) preserved. Beverly Burlingame (collector), 12.vi.2021.

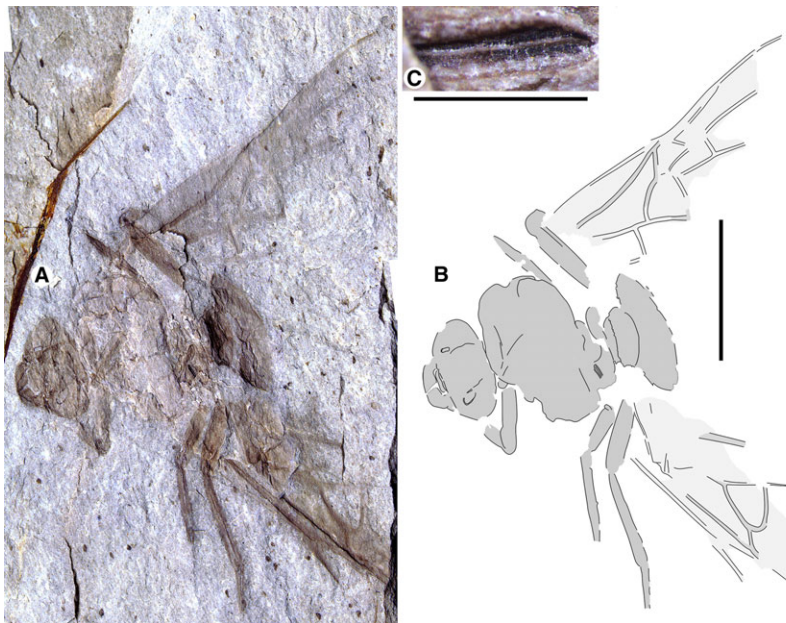


Fig. 1. Allenby Formation ant, *Titanomyrma* sp. UBC-BBM-PAL-2022-00001. **A**, photograph; **B**, drawing; and **C**, partially preserved spiracle. **A**, **B** to scale = 1 cm; **C** to scale = 1 mm.

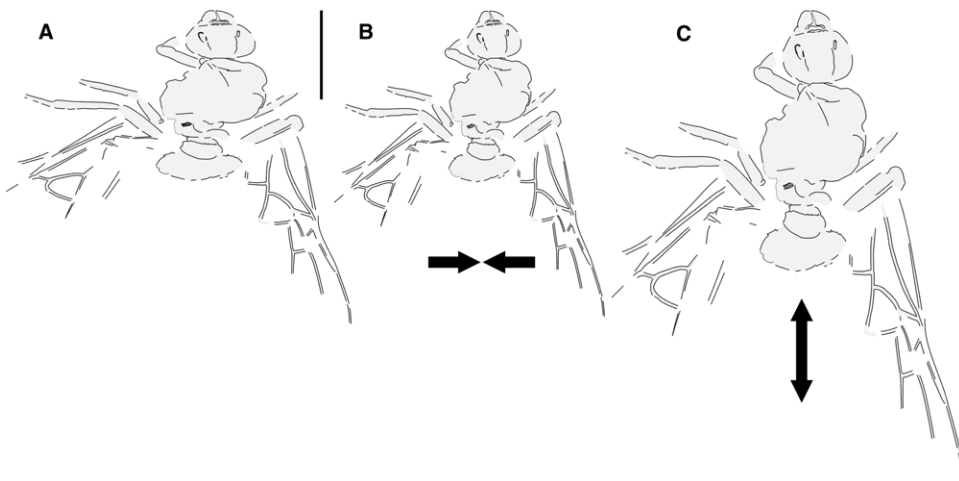
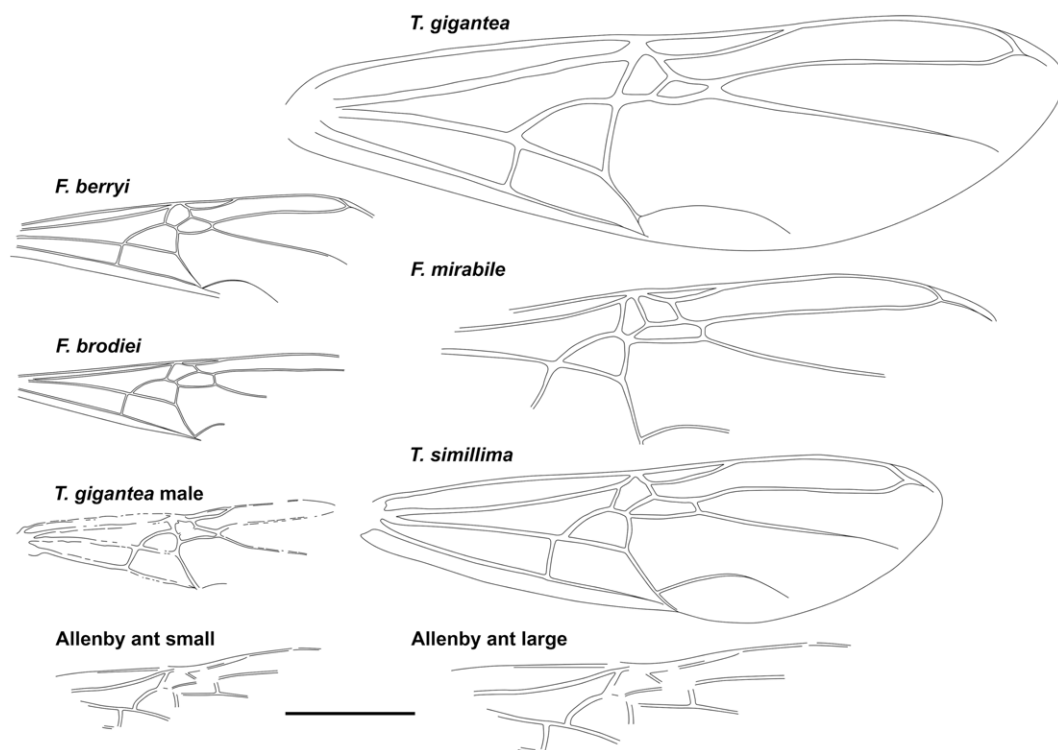


Fig. 2. Allenby Formation ant *Titanomyrma* sp. BBM-PAL-2022-00001: **A**, as preserved; **B**, compressed laterally; and **C**, extended lengthwise (see text).

**The Green River Formation ants.** Fossil Butte National Monument collection FOBU-9488A, B (part and counterpart), queen (Fig. 4).

Head: missing antennae; head length, ~ 5.1 mm; head width, ~ 6.5 mm; eye length, ~ 1.1 mm, head capsule about 2/5 alitrunk length (alitrunk indistinctly preserved), with rounded posterior margins, not lobed; eyes large, oval, maximum length oriented anterodorsally, located at cephalic mid-length; large, triangular mandibles, about half head length, with approximately 8–10 coarse teeth. Alitrunk: somewhat distorted, indistinctly preserved; wings not preserved; right legs



**Fig. 3.** Comparative forewings of species indicated with the Allenby ant compressed and lengthened. The wings of *F. berryi*, *F. brodiei*, *T. gigantea* (queen), and *T. simillima* were redrawn from Lutz (1986), figs. 2C, 1C, 5, and 10, respectively, and the wing of the *T. gigantea* male was redrawn from Wappler (2003), fig. 94. Scale = 1 cm.

partially preserved, short, stout, left legs absent. Spiracles slit-like, long; propodeal spiracles robust, widest; petiole spiracles smaller, thinner; one gaster spiracle (A3) present, poorly preserved. Waist: single segmented, petiole lacking anterior peduncle, apparently about 5 mm wide, otherwise indistinct. Gaster: small, indistinct portion preserved. Paul Buchheim (collector), 7.viii.2002.

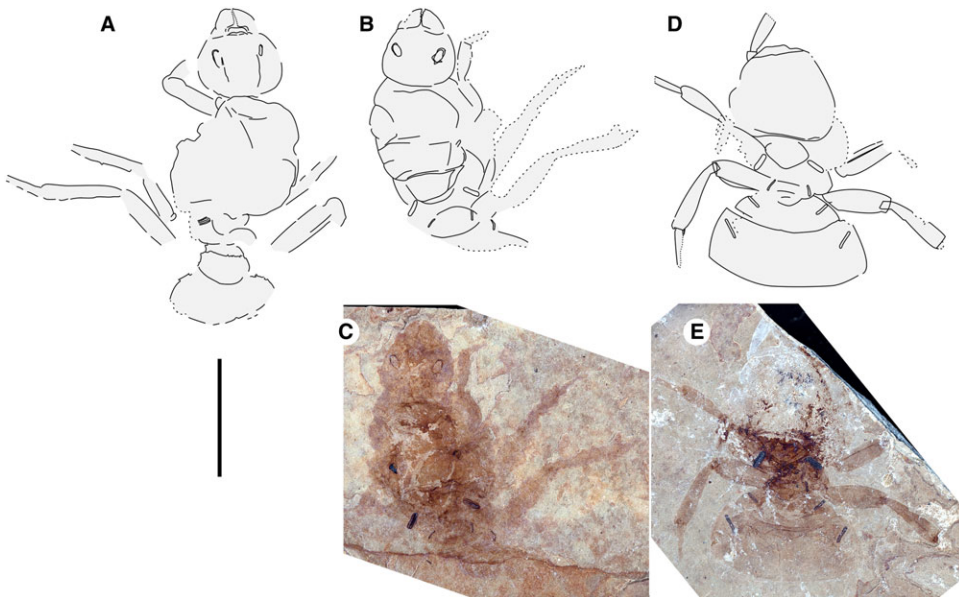
Fossil Butte National Monument collection FOBU-6495 (part only), queen (Fig. 4).

Head missing. Alitrunk: well preserved basally, faintly in anterior two thirds; legs short, stout; femur, tibia of the left mesothoracic and metathoracic legs, the right metathoracic leg, the femur of right mesothoracic leg, and tibia of the left prothoracic present. Wings not preserved. Spiracles slit-like, long; propodeal spiracle robust, widest; petiole spiracle shortest; petiole, A3, A4 spiracles narrower; A3, A4 spiracles long. Waist: single segmented, petiole about 4.5 mm wide, length indeterminate, apparently lacking anterior peduncle. Gaster: A3, about 9 mm wide, length indeterminate; A4, 14 mm wide, about 4.5–5.0 mm long, missing posterior to this. Paul Buchheim (collector), 25.vii.2000.

## Discussion

### Ant size, climate, and trans-Arctic dispersal

The new Allenby specimen is the first occurrence of Formiciinae in a known microthermal climate. Its implications for the relationships between queen size, climate, and trans-Arctic dispersal are complicated, however, by uncertainty of its size in life, as its sole fossil has been



**Fig. 4.** **A**, Allenby Formation ant, as lengthened (compare with Fig. 2C), and new Green River Formation ants – *Titanomyrma* sp.: **B**, FOBU 9488B drawing; **C**, FOBU 9488B photograph; **D**, FOBU 6495 drawing; and **E**, FOBU 6495 photograph. All to scale, 1 cm.

distorted by geologic shear forces during diagenesis, a common occurrence in the exposure where it was found (see Archibald and Cannings 2021, fig. 3, a similarly distorted fossil from the same locality). It was either compressed lengthwise to shorten it or stretched at a right angle from this to widen it (Fig. 2). These possibilities cannot be distinguished.

**The Allenby ant as big.** Graphically stretching the fossil a few degrees from lengthwise to be bilaterally symmetrical (Fig. 2C) results in a shape and size indistinguishable from the undistorted and more clearly preserved but incomplete new Green River specimens (Fig. 4) and closely resembles the sizes of *T. lubei* and *T. simillima* queens, both about 5 cm long (wings, Fig. 3).

That modern ant species with the largest queens inhabit low latitudes (Archibald *et al.* 2011b) seems to contradict the finding that the average body mass of worker ants decreases along with increasing mean annual temperatures towards the equator (Kaspari 2009). This might be accounted for by increasing ant species richness towards the equator, lengthening the upper tail of size distribution of workers; that is, although average worker size decreases in low latitudes, individual species with the largest workers are found there – for example, those of *Dinoponera* Roger, *Paraponera* Smith, and some *Pachycondyla* Smith. If this relationship holds true for queens, a giant Allenby queen could result from the high, tropical levels of species richness of insects in the Okanagan Highlands, associated with seasonal temperature stability (Archibald *et al.* 2010). Whatever explains gigantism in *Titanomyrma* queens, the Allenby species being giant would indicate that huge size is decoupled from mean annual temperatures, which would support the hypothesis that the genus was cyrophobic. This would allow giant species to ignore hyperthermal gates and migrate across high latitudes at any time when the land bridges existed, including times of cooler mean annual temperature values. Other truly thermophilic organisms would still require hyperthermals to cross the Arctic, even if the giant formiciniines did not.

**The Allenby ant as small.** Graphically compressing the fossil laterally (Fig. 2B) produces an identical but smaller shape, estimated to be about 3.3 cm in length presumed complete, approximately 65% the size of the stretched version and of the *T. simillima* and *T. lubei* queens. Its forewings are then comparable in size to those of the smaller formiciniines *F. brodiei* and *F. berryi* or to males of *T. simillima* and *T. gigantea* (Fig. 3). It has been suggested that specimens of *F. brodiei* and *F. berryi*, known only by forewings, might have been males (Fig. 3; Lutz 1986; Katzke *et al.* 2017). The Allenby ant is not a male, because these are most easily distinguished from queens in *T. simillima* and *T. gigantea* by their notably smaller relative head size. Such a small Allenby formiciniine queen, although bigger than the 3-cm size set as large for modern ants in Archibald *et al.*'s (2011b) analysis, would support the hypothesis that species of Formiciniine with the largest queens required high mean annual temperatures, whereas smaller species inhabited a wide variety of climates, including cooler microthermal ones that excluded species with large queens. This would be consistent with the hypothesis that giant Formiciniinae crossed the Arctic only during hyperthermals and smaller species crossed at any time the corridors were open.

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## References

- Archibald, S.B. and Cannings, R.A. 2021. The head of Cephalozygoptera (Odonata). *Zootaxa*, **5047**: 97–100. <https://doi.org/10.11646/zootaxa.5047.1.10>.
- Archibald, S.B. and Farrell, B.D. 2003. Wheeler's dilemma. *Proceedings of the Second Paleontomological Congress, Acta Zoologica Cracoviensia*, **46** (supplement, Fossil Insects): 17–23.
- Archibald, S.B., Bossert, W.H., Greenwood, D.R., and Farrell, B.D. 2010. Seasonality, the latitudinal gradient of diversity, and Eocene insects. *Paleobiology*, **36**: 374–398. <https://doi.org/10.1666/09021.1>
- Archibald, S.B., Greenwood, D.R., Smith, R.Y., Mathewes, R.W., and Basinger, J.F. 2011a. Great Canadian Lagerstätten 1. Early Eocene Lagerstätten of the Okanagan Highlands (British Columbia and Washington State). *Geoscience Canada*, **38**: 155–164. Available from <https://journals.lib.unb.ca/index.php/GC/article/view/18964> [accessed 3 December 2022].
- Archibald, S.B., Johnson, K.R., Mathewes, R.W., and Greenwood, D.R. 2011b. Intercontinental dispersal of giant thermophilic ants across the Arctic during early Eocene hyperthermals. *Proceedings of the Royal Society B*, **278**: 3679–3686. <https://doi.org/10.1098/rspb.2011.0729>.



- Archibald, S.B., Morse, G.E., Greenwood, D.R., and Mathewes, R.W. 2014. Fossil palm beetles refine upland winter temperatures in the Early Eocene Climatic Optimum. *Proceedings of the National Academy of Sciences of the United States of America*, **111**: 8095–8100. <https://doi.org/10.1073/pnas.1323269111>.
- Brikiatis, L. 2014. The De Geer, Thulean and Beringia routes: key concepts for understanding early Cenozoic biogeography. *Journal of Biogeography*, **41**: 1036–1054. <https://doi.org/10.1111/jbi.12310>.
- Buchheim, H.P., Cushman, R.A., and Biaggi, R.E. 2011. Stratigraphic revision of the Green River Formation in Fossil Basin, Wyoming: overfilled to underfilled lake evolution. *Rocky Mountain Geology*, **46**: 165–181. <https://doi.org/10.2113/gsrocky.46.2.165>.
- Eberle, J.J. and Greenwood, D.R. 2012. Life at the top of the greenhouse Eocene world: a review of the Eocene flora and vertebrate fauna from Canada's High Arctic. *Geological Society of America Bulletin*, **124**: 3–23. <https://doi.org/10.1130/B30571.1>.
- Erwin, D.W. and Stockey, R.A. 1991. Silicified monocotyledons from the Middle Eocene Princeton chert (Allenby Formation) of British Columbia, Canada. *Review of Palaeobotany and Palynology*, **70**: 147–162. [https://doi.org/10.1016/0034-6667\(91\)90083-F](https://doi.org/10.1016/0034-6667(91)90083-F).
- Greenwood, D.R., Archibald, S.B., Mathewes, R.W., and Moss, P.T. 2005. Fossil biotas from the Okanagan Highlands, southern British Columbia and northern Washington State: climates and ecosystems across an Eocene landscape. *Canadian Journal of Earth Sciences*, **42**: 167–185. <https://doi.org/10.1139/E04-100>.
- Kaspari, M. 2009. Global energy gradients and size in colonial organisms: worker mass and worker number in ant colonies. *Proceedings of the National Academy of Sciences of the United States of America*, **102**: 5079–5083. <https://doi.org/10.1073/pnas.0407827102>.
- Katzke, J., Barden, P., Dehon, M., Michez, D., and Wappler, T. 2017. Giant ants and their shape: revealing relationships in the genus *Titanomyrma* with geometric morphometrics. *PeerJ*, **6**: e4242. <https://doi.org/10.7717/peerj.4242>.
- Kim, J.-H., van der Meer, J., Schouten, S., Helmke, P., Willmott, V., Sangiorgi, F., *et al.* 2010. New indices and calibrations derived from the distribution of crenarchaeal isoprenoid tetraether lipids: implications for past sea surface temperature reconstructions. *Geochimica et Cosmochimica Acta*, **74**: 4639–4654. <https://doi.org/10.1016/j.gca.2010.05.027>.
- Lutz, H. 1986. Eine neue Unterfamilie der Formicidae (Insecta: Hymenoptera) aus dem mittelozeänen Ölschiefer der 'Grube Messel' bei Darmstadt (Deutschland, S-Hessen) [A new subfamily of Formicidae (Insecta: Hymenoptera) from the Middle Eocene oil shale of the 'Grube Messel' near Darmstadt (Germany, S-Hesse)]. *Senckenbergiana Lethaea*, **67**: 177–218.
- Lutz, H. 1990. Systematische und palökologische Untersuchungen an Insekten aus dem Mittel-Eozän der Grube Messel bei Darmstadt [Systematic and palaeoecological studies on insects from the Middle Eocene of the Messel Pit near Darmstadt]. *Courier Forschungsinstitut Senckenberg*, **124**: 1–165.
- Mathewes, R., Greenwood, D.R., and Archibald, S.B. 2016. Paleoenvironment of the Quilchena flora, British Columbia, during the Early Eocene Climatic Optimum. *Canadian Journal of Earth Sciences*, **53**: 574–590. <https://doi.org/10.1139/cjes-2015-0163>.
- McInerney, F.A. and Wing, S.L. 2011. The Paleocene–Eocene thermal maximum: a perturbation of carbon cycle, climate, and biosphere with implications for the future. *Annual Review of Earth and Planetary Sciences*, **39**: 489–516. <https://doi.org/10.1146/annurev-earth-040610-133431>.
- National Research Council. 2011. *Understanding Earth's deep past: lessons for our climate future*. The National Academies Press, Washington, District of Columbia, United States of America. 194 pp. <https://doi.org/10.17226/13111>.

- Read, P. 2000. Geology and industrial minerals of the Tertiary basins, south–central British Columbia. GeoFile 2000-3. Geological Survey, British Columbia Ministry of Energy and Mines, Victoria, British Columbia, Canada. 110 pp. Available from [https://cmscontent.nrs.gov.bc.ca/geoscience/publicationcatalogue/GeoFile/BCGS\\_GF2000-03.pdf](https://cmscontent.nrs.gov.bc.ca/geoscience/publicationcatalogue/GeoFile/BCGS_GF2000-03.pdf) [accessed 19 November 2022].
- Reinhardt, L., von Gosen, W., Lückge, A., Blumenberg, M., Galloway, J.M., West, C.K., *et al.* 2022. Geochemical indications for the Paleocene–Eocene thermal maximum (PETM) and Eocene thermal maximum 2 (ETM-2) hyperthermals in terrestrial sediments of the Canadian Arctic. *Geosphere*, **18**: 327–349. <https://doi.org/10.1130/GES02398.1>.
- Rubino, E., Leier, A., Cassell, E.J., Archibald, S.B., Foster-Baril, Z., and Barbeau, D.L. 2021. Detrital zircon U–Pb ages and Hf-isotopes from Eocene intermontane basin deposits of the southern Canadian Cordillera. *Sedimentary Geology*, **422**: 105969. <https://doi.org/10.1016/j.sedgeo.2021.105969>.
- Sluijs, A., Frieling, J., Inglis, G.N., Nierop, K.G.J., Peterse, F., Sangiorgi, F., and Schouten, S. 2020. Late Paleocene–early Eocene Arctic Ocean sea surface temperatures: reassessing biomarker paleothermometry at Lomonosov Ridge. *Climate of the Past*, **16**: 2381–2400. <https://doi.org/10.5194/cp-16-2381-2020>.
- Smith, M.E. and Carroll, A.R. 2015. Introduction to the Green River Formation. *In* Stratigraphy and paleolimnology of the Green River Formation, western USA. Syntheses in Limnogeology. Volume 1. *Edited by* M.E. Smith and A.R. Carroll. Springer, New York, New York, United States of America. Pp. 1–12. [https://doi.org/10.1007/978-94-017-9906-5\\_1](https://doi.org/10.1007/978-94-017-9906-5_1).
- Stokke, E.W., Jones, M.T., Tierney, J.E., Svensen, H.H., and Whiteside, J.H. 2020. Temperature changes across the Paleocene–Eocene thermal maximum: a new high-resolution TEX86 temperature record from the Eastern North Sea Basin. *Earth and Planetary Science Letters*, **544**: 116388. <https://doi.org/10.1016/j.epsl.2020.116388>.
- Sunderlin, D., Loope, G., Parker, N.E., and Williams, C.J. 2011. Paleoclimate and paleoecological implications of a Paleocene–Eocene fossil leaf assemblage, Chickaloon Formation, Alaska. *Palaios*, **26**: 335–345. <https://doi.org/10.2110/palo.2010.p10-077r>.
- Vickers, M.L., Lengger, S.K., Bernasconi, S.M., Thibault, N., Schultz, B.P., Fernandez, A., *et al.* 2020. Cold spells in the Nordic Seas during the early Eocene Greenhouse. *Nature Communications*, **11**: 4713. <https://doi.org/10.1038/s41467-020-18558-7>.
- Wappler, T. 2003. Die Insekten aus dem Mittel-Eozän des Eckfelder Maars, Vulkaneifel [Insects from the Middle Eocene of the Eckfeld Maar, Vulkaneifel]. *Mainzer Naturwissenschaftliches Archiv*, **27**: 1–234.
- West, C.K., Greenwood, D.R., Reichgelt, T., Lowe, A.J., Vachon, J.M., and Basinger, J.F. 2020. Paleobotanical proxies for early Eocene climates and ecosystems in northern North America from middle to high latitudes. *Climate of the Past*, **16**: 1387–1410. <https://doi.org/10.5194/cp-16-1387-2020>.
- Westerhold, T., Röhl, U., Frederichs, T., Agnini, C., Raffi, I., Zachos, J.C., and Wilkens, R.H. 2017. Astronomical calibration of the Ypresian timescale: implications for seafloor spreading rates and the chaotic behavior of the solar system? *Climate of the Past*, **13**: 1129–1152. <https://doi.org/10.5194/cp-13-1129-2017/>
- Wilf, P. 2000. Late Paleocene–early Eocene climate changes in southwestern Wyoming: paleobotanical analysis. *Geological Society of America Bulletin*, **112**: 292–307. [https://doi.org/10.1130/0016-7606\(2000\)112<292:LPECCI>2.0.CO;2](https://doi.org/10.1130/0016-7606(2000)112<292:LPECCI>2.0.CO;2).
- Wing, S.L. and Greenwood, D.R. 1993. Fossils and fossil climate: the case for equable continental interiors in the Eocene. *In* Palaeoclimates and their modelling, with special reference to the Mesozoic Era. *Edited by* J.R.L. Allen, B.J. Hoskins, B.W. Sellwood, and R.A. Spicer. *Philosophical Transactions of the Royal Society of London B*, **341**: 243–252. <https://doi.org/10.1098/rstb.1993.0109>.

- Wing, S.L., Harrington, G.J., Smith, F.A., Bloch, J.I., Boyer, D.M., and Freeman, K.H. 2005. Transient floral change and rapid global warming at the Paleocene–Eocene boundary. *Science*, **310**: 993–996. <https://doi.org/10.1126/science.1116913>.
- Wolfe, J.A. 1975. Some aspects of plant geography of the Northern Hemisphere during the Late Cretaceous and Tertiary. *Annals of the Missouri Botanical Garden*, **62**: 264–279. <https://doi.org/10.2307/2395198>.
- Zachos, J.C., Dickens, G.R., and Zeebe, R.E. 2008. An early Cenozoic perspective on greenhouse warming and carbon-cycle dynamics. *Nature*, **451**: 279–283. <https://doi.org/10.1038/nature06588>.
- Zachos, J.C., McCarren, H., Murphy, B., Röhl, U., and Westerhold, T. 2010. Tempo and scale of late Paleocene and early Eocene carbon isotope cycles: implications for the origin of hyperthermals. *Earth and Planetary Science Letters*, **299**: 242–249. <https://doi.org/10.1016/j.epsl.2010.09.004>.

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