Holocene deglaciation and climate history of the northern Antarctic Peninsula region: a discussion of correlations between the Southern and Northern Hemispheres

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ABSTRACT. The chronology of post-Last Glacial Maximum deglaciation in the northern Antarctic Peninsula region is discussed. It is concluded that, contrary to what was earlier believed, the deglaciation process here was largely out-of-phase with that in the Northern Hemisphere. Although, for global eustatic reasons, the marine-based glaciers may have retreated simultaneously with ice-melting in the Northern Hemisphere, the land-based glaciers retreated only slowly during the first half of the Holocene, about 9000–5000 BP. This may have been due either to increased precipitation counterweighing ablation or to delayed warming. A distinct but rather brief glacial readvance took place around 5000 BP, probably caused by a period of renewed cooling. It was followed by the Holocene climatic optimum, about 4000–3000 BP. This warm "hypsithermal" period thus came much later than its equivalent in the Northern Hemisphere, but it roughly coincided with the Milankovitchean Holocene insolation maximum for these southern latitudes.

INTRODUCTION

In the early days of Quaternary studies in the Antarctic Peninsula region (including the South Shetland Islands and the islands in the western Weddell Sea), it was generally believed that the post-Last Glacial Maximum (Wisconsinan, Weichselian, etc.) deglaciation had a timing similar to that in the Northern Hemisphere. The onset of deglaciation of the presently ice-free areas was assumed to have begun before 10 000 BP, and the oldest isostatically elevated sediments related to the deglaciation were dated to about 9000 BP (e.g. John and Sugden, 1971; Sugden and John, 1973). These ideas prevailed until the beginning of this decade (e.g. Clapperton, 1990), although hints of a more complicated picture had been published several years before (Rabassa, 1983, 1987).

Work carried out on land in the South Shetlands and on the islands in the western Weddell Sea during the 1980s and 1990s suggests a different deglaciation pattern than "the Northern Hemisphere concept". Studies of glacial deposits and their relation to isostatically elevated marine sediments, and of lake-sediment and moss-bank cores, suggest that (1) the initial deglaciation of the presently ice-free lowland areas in the Antarctic Peninsula region took place gradually between about 9000 and 5000 14C years BP, and (2) that a distinct but rather short-lived glacial readvance culminated around or shortly after 5000 BP, followed (3) by the Holocene climatic optimum ("hypsithermal") between about 4000 and 3000 BP. Thereafter, (4) climate gradually approached present conditions, as illustrated by a number of mostly restricted neoglacial oscillations, including a correlative of the Northern Hemisphere's Little Ice Age readvance.

It also seems that most of the data which were earlier interpreted according to the "Northern Hemisphere concept" may be fitted into the new paradigm. This is especially the case when the empirically derived marine reservoir-effect correction of -1200 years (e.g. Gordon and Harkness, 1992) is applied, instead of the smaller corrections (700–850 years) used in earlier works.

The present paper gives a brief overview of the available terrestrial dataset, and tries to put some of the "older" data into the modern concept.

Table 1. Radiocarbon datings regarded as closely postdating the last deglaciation of different areas in the northern Antarctic Peninsula region (Fig. 1). The dates on marine organisms are corrected for a reservoir effect of -1200 years (Gordon and Harkness, 1992)

Locality	Age BP	Dated sample	Source
Jurasee, Fildes Peninsula, King George Island	8700 ± 300	Limnic moss?	Mäusbacher (1991)
Potter's Cove King George Island	8470 ± 230	Mollusc frag- ments	Sugden and John (1973)
The Naze James Ross Island	7360 ± 100	Mollusc frag- ments	Hjort and others (1997)
The Naze James Ross Island	7260 ± 90	Mollusc shells	Hjort and others (1997)
Walker Point, Elephant Island	5350 ± 60	Moss peat	Björck and others (1991)
Lake 35/"Åsa", Byers Peninsula, Livingston Island	4600 ± 100	Limnic moss	Björck and others (1993, 1996a)

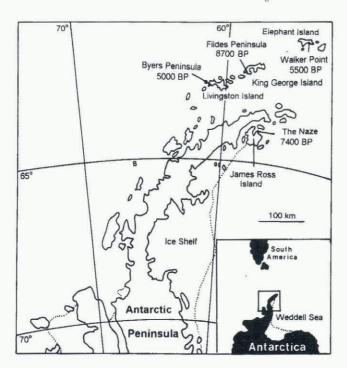


Fig. 1. Map of the northern Antarctic Peninsula region, showing ¹⁴C dates regarded as closely postdating the deglaciation in different areas.

THE DEGLACIATION

Recently published land-based studies of the post-Last Glacial Maximum deglaciation in the Antarctic Peninsula region are from three main areas: (1) Fildes Peninsula on King George Island (e.g. Mäusbacher, 1991), (2) Byers Peninsula on Livingston Island (e.g. Björck and others, 1996a), both in the South Shetland Islands; and (3) James Ross Island in the western Weddell Sea (Ingólfsson and others, 1992; extended and updated by Hjort and others, 1997). The oldest reliable ¹⁴C datings from the region (see under James Ross Island, below) are listed in Table 1 and indicated in Figure 1.

Fildes Peninsula

The Fildes Peninsula studies (Mäusbacher, 1991, e.g. fig. 72) are based on ¹⁴C-dated lake-sediment cores. They indicate a slow and gradual deglaciation of this lowland area (the second largest in the South Shetlands, after Byers Peninsula) between about 9000 BP (the oldest dating is about 8700 BP; see Table 1) and 5500 BP. During that time, sea level never stood above 22 m, and it took the ice front more than 3000 years to retreat 5 km! Around 5000 BP the minerogenic component in the lake sediments distinctly increased. This is interpreted as the result of renewed glacial activity, which lasted until about 4000 BP.

Sugden and John's (1973) oldest deglaciation date from the South Shetlands, from nearby Potter's Cove on King George Island (detailed description in John and Sugden, 1971), was taken as evidence for a deglaciation around 9000 BP. However, when using a -1200 years reservoir correction, that date becomes only about 8500 BP (Table 1).

Mäusbacher (1991) also dated a distinct 16 m shoreline on Fildes Peninsula (part of the 15–20 m main Holocene beach level in the region) to about 6000 BP, which roughly fits with the approximately 5400 BP age of the 18 m level in Potter's Cove obtained when using the -1200 years correction on Clapperton and Sugden's (1988) original date.

Byers Peninsula

The oldest deglaciation date from Byers Peninsula, about 4600 BP (Björck and others, 1993, 1996a), is also from lake sediments. The cores from this area show that the "postglacial" marine limit here lies below 35 m. Deglaciation proceeded gradually until most of the approximately 6 km wide (in the direction the glacier withdrew) peninsula was ice-free by about 3000 BP, and the retreat was faster than on Fildes Peninsula during the earlier part of the Holocene. The ice front eventually withdrew to somewhere behind its present limit, as shown by a whale bone transported forwards and upwards during the Little Ice Age readvance of the Rotch Dome ice front, and by beach ridges continuing underneath the ice cap (Björck and others, 1996a).

James Ross Island

A recent critical discussion of deglaciation dates from James Ross Island and surrounding islands in the western Weddell Sea (Hjort and others, 1997) casts doubt on the reliability of two ¹⁴C dates with ages between 10 000 and 9000 BP. These were made on algal mats and moss from glacier-margin lakes (Zale and Karlén, 1989; Ingólfsson and others, 1992) and may have been contaminated by old carbon from the glacial meltwater. It is suggested (Hjort and others, 1997) that deglaciation of the presently ice-free lowland areas here started only shortly before 7400 BP, at a marine level around 30 m. The mid-Holocene "Bahia Bonita" glacial readvance, first described by Rabassa (1983) from Brandy Bay on northern James Ross Island and then 14C-dated to about 5000 BP, was now found to have culminated at around 4600 BP. This was after a non-surge advance of more than 7 km (Hjort and others, 1997). The advance was contemporaneous with a marine level at 18-16 m. By about 4300 BP the lowlands in Brandy Bay were ice-free again, and from studies of lake sediments Björck and others (1996b) dated the Holocene climatic optimum (a "hypsithermal" period with relatively very warm and humid conditions) to about 4200-3000 BP.

A glacial oscillation marked by a set of distinct moraines at Hope Bay, near the tip of the Antarctic Peninsula north of James Ross Island, has been dated to about 4700 BP by Zale (1994).

DISCUSSION AND CONCLUSIONS

It has not yet been established exactly how far out from the present coastline the glaciers around the Antarctic Peninsula and its surrounding islands reached during the Last Glacial Maximum, partly as a dynamic effect of the lowering of the grounding line with the global eustatic sea-level regression. It can, however, be postulated that when global sea level started to rise again, from about 15 000 BP onwards, these glacier fronts gradually retreated (e.g. Pope and Anderson, 1992; Pudsey and others, 1994).

Thus far, parallelism with the Northern Hemisphere is uncontroversial. However, when it comes to the chronology of the deglaciation of the presently ice-free lowland areas, which to a large extent remained above sea level throughout the Holocene, the picture differs much from that in the Northern Hemisphere. It also differs from most Antarctic ice-core data interpretations, which have tended to see a strong parallelism with the Greenland ice-core data. In the areas of Antarctica treated in this paper, there seems to have been a gradual, often slow deglaciation, starting at the earliest around 9000 BP and with its first phase not completed until about 5500 BP (Björck and others, 1991; Mäusbacher, 1991). Then came the "Bahia Bonita" glacial readvance, culminating shortly after 5000 BP. This advance was followed by renewed glacial retreat, more rapid in the continental areas of higher equilibrium-line altitude (ELA) in the western Weddell Sea (like James Ross Island) than in the maritime, lower-ELA South Shetlands (such as on Byers Peninsula). Then, about 4000–3000 BP, came the short Antarctic Peninsula region hypsithermal, after which conditions gradually approached the present.

The extended period of rather slow deglaciation during the first half of the Holocene (not unlike what has recently been suggested for the inner parts of the Ross Sea; e.g. Licht and others, 1996) could either indicate a mass-balance situation with increased precipitation during the gradual establishment of Holocene interglacial conditions or simply be due to delayed (as compared with the Northern Hemisphere) warming. The glacial readvance culminating around 5000–4500 BP (also identified along the East Antarctic coast, by Domack and others, 1991) now seems most likely to have been an effect of renewed cooling. Such a colder period, lasting from about 7000 to 4500 BP, has recently been identified by Shevenell and others (1996) from marine sediments on the western side of the Antarctic Peninsula

The final disappearance of glaciers from the lowland areas discussed roughly coincided with the beginning of the brief hypsithermal, which as noted lasted between about 4000 and 3000 BP and was the warmest Holocene period on record in this region (e.g. Björck and others, 1996b; Shevenell and others, 1996). This warm period has also been recorded in the Ross Sea area, where it was expressed by expanding penguin colonisation (Baroni and Orombelli, 1994). It approximately coincided with the peaking of Milankovitchean summer insolation in the south (e.g. Budd and Smith, 1987; Budd and Rayner, 1990), which, especially in the northern Antarctic Peninsula region, was well out of phase with insolation in the Northern Hemisphere.

We thus conclude that the geological ground truth in the Antarctic Peninsula region does not suggest any close chronological correlation with Holocene glacial events in the Northern Hemisphere. Except for the purely eustatically induced changes, glacial developments in this region seem to have been largely out-of-phase with events in the north. An exception was the Little Ice Age.

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