

## EXTRACTION AND AMS RADIOCARBON DATING OF POLLEN FROM LAKE BAIKAL SEDIMENTS

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**ABSTRACT.** This work focuses on the preparation and dating of sporomorph (pollen and spores) concentrates of high purity. Three sediment cores recovered from Lake Baikal within the EU-Project CONTINENT were subjected to palynological analyses and accelerator mass spectrometry (AMS) radiocarbon dating. Laboratory processing of concentrates was aimed at the removal of non-sporomorph organic matter by means of chemical treatment, micro-sieving, and heavy liquid separation. The obtained concentrates were checked under the microscope and sample purity was estimated on the basis of particle counts. The results of AMS <sup>14</sup>C dating show differences in the sedimentation rate among 3 sites of Lake Baikal.

### METHODS

Lake Baikal is situated in southeast Siberia in the eastern part of the Russian Federation near the Mongolian border. It is the largest (20% of the total volume of surface freshwater), deepest (>1600 m), and probably oldest (about 25 million yr) lake on Earth.

Radiocarbon dating of Lake Baikal sediments is a difficult challenge, as previous studies have proved (Coleman 1996). The main problem is the scarcity of material suitable for dating because the sediments are very poor in organic matter and carbonates. Our work focused on the preparation and dating of high-purity sporomorph concentrates.

Three sediment cores recovered from Lake Baikal were subjected to palynological analyses and AMS <sup>14</sup>C dating. The cores come from 3 different locations: Posolskoe Bank, Vydrino Shoulder, and Continent Ridge (see Figure 1). The cores subjected to dating were collected with the use of a Kasten corer, which gives the best assurance of complete recovery of the sediment.

The preliminary knowledge about sporomorph concentrations and sporomorph assemblages in Holocene and Late Glacial core sections, necessary for selecting appropriate subsampling levels, was gained during palynological analyses of these cores. Sample sizes ranged between 20 and 160 cm<sup>3</sup> (see Table 1), depending on available material, pollen concentrations, and a rough calculation of pollen mass and carbon content using data available from previous studies (Erdtman 1969; Brown et al. 1989). The volume of samples taken for preparation was adjusted to obtain a minimum carbon content of about 2–4 mg per sample.

### Extraction of Sporomorphs

Laboratory processing of pollen and spore concentrates was aimed at achieving a high purity of the samples. The removal of inorganic and non-sporomorph organic matter was carried out using chemical treatment, micro-sieving, and heavy liquid separation.

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Table 1 Details concerning preparation of pollen concentrates and graphite targets and results of dating.

Sample code	Sample name	Depth range (cm)	Sample volume (cm <sup>3</sup> )	Estimated purity	Carbon yield (%)	Target (mg C)	Age BP (yr)	Age cal BP (yr)	$\delta^{13}\text{C}$ (‰)
<b>Výdrino Shoulder</b>									
1	GdA-260	0–4	60	90%	73	1.00	1950 ± 35	1820–1990	-27.2 ± 0.5
2	GdA-261	20–24	80	95%	73	1.01	2375 ± 25	2330–2710	-26.1 ± 0.4
3	GdA-263	32–35	60	95%	58	1.01	2945 ± 30	2990–3220	-26.1 ± 0.7
4	GdA-264	56–59	60	95%	62	1.01	4530 ± 40	5040–5320	-24.7 ± 0.7
5	GdA-265	77–80	60	95%	65	1.01	5455 ± 35	6170–6310	-21.5 ± 0.5
6	GdA-266	89–92	60	90%	65	1.00	6170 ± 40	6940–7230	-23.9 ± 1.1
7	GdA-267	101–104	60	85%	75	1.01	6700 ± 40	7480–7670	-20.9 ± 0.5
8	GdA-268	121–124	60	85%	79	1.01	7760 ± 40	8420–8600	-22.0 ± 0.6
9	GdA-269	138–141	60	90%	66	1.01	8620 ± 40	9520–9700	-20.1 ± 0.9
10	GdA-271	145–148	60	90%	71	1.00	8750 ± 50	9550–9950	-26.2 ± 0.7
11	GdA-272	159–163	80	90%	69	1.00	9470 ± 50	10,550–11,100	-23.0 ± 0.5
12	GdA-270	171.5–175.5	80	80%	62	1.00	10,030 ± 50	11,250–11,950	-22.3 ± 0.3
<b>Continent Ridge</b>									
1	GdA-247	0–4	100	70%	89	0.99	1235 ± 25	1060–1270	-25.2 ± 0.3
2	GdA-248	18–22	80	60%	75	0.99	4060 ± 30	4420–4790	-22.8 ± 0.3
3	GdA-249	24–28	80	70%	77	1.00	4955 ± 30	5600–5740	-25.0 ± 0.2
4	GdA-250	32–36	80	85%	73	1.00	5335 ± 30	5990–6270	-24.5 ± 0.2
5	GdA-251	39–43	100	70%	53	0.99	6630 ± 35	7430–7580	-26.2 ± 0.4
6	GdA-252	54–58	100	80%	69	1.01	8570 ± 40	9470–9630	-25.3 ± 0.3
7	GdA-253	60–64	80	90%	40	1.00	8690 ± 50	9540–9890	-28.7 ± 0.9
8	GdA-257	70–76	120	85%	43	1.01	10,250 ± 50	11,650–12,400	-24.2 ± 0.8
9	GdA-258	78–84	120	10%	80	0.27	11,090 ± 80	12,650–13,400	-47.0 ± 2.1
10	GdA-259	86–92	120	60%	59	0.82	11,960 ± 50	13,550–15,250	-26.5 ± 0.6
<b>Posolskoe Bank</b>									
1	GdA-217	29–32	30	65%	62	1.00	2425 ± 40	2340–2720	-35.2 ± 0.6
2	GdA-219	44–46	30	70%	48	1.01	3200 ± 45	3330–3560	-29.3 ± 0.7
3	GdA-231	56–57	20	85%	65	0.99	3850 ± 35	4150–4410	-25.9 ± 0.1
4	GdA-232	57–58	20	90%	64	0.99	3985 ± 40	4300–4570	-30.0 ± 0.8
5	GdA-220	66–68	30	85%	40	0.99	5520 ± 45	6200–6410	-28.8 ± 0.4
6	GdA-225	80–81	20	60%	60	0.99	9350 ± 60	10,280–10,740	-28.3 ± 0.4
7	GdA-226	81–82	20	60%	73	0.93	11,710 ± 70	13,300–13,800	-34.4 ± 0.7
8	GdA-233	84–86	40	70%	60	0.25	11,320 ± 110	13,300–13,900	-44.9 ± 0.4
9	GdA-218	89–93	60	<80%	84	0.71	12,200 ± 70	13,750–15,450	-29.6 ± 0.6
10	GdA-224	100–104	60	<30%	19	0.16	15,250 ± 210	17,410–18,370	-52.3 ± 2.6
11	GdA-223	108–112	80	<60%	65	0.93	15,040 ± 90	17,360–18,040	-27.4 ± 0.6
12	GdA-222	116–124	160	<60%	65	0.85	20,920 ± 160	23,760–25,040	-33.5 ± 0.8

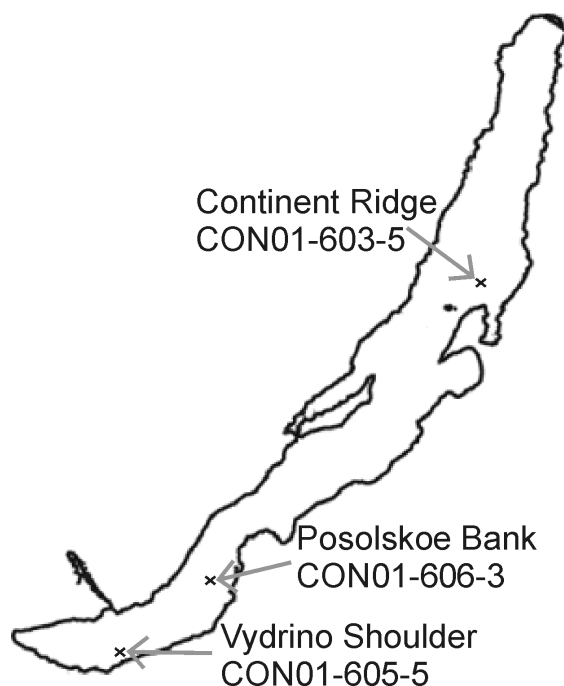


Figure 1 Map showing the location of sites: Continent Ridge (52.95460°N, 108.91370°E; water depth 130 m; collected 29 Jul 2001); Posolskoe Bank (52.08390°N, 105.86540°E; water depth 386 m; collected 3 Aug 2001); and Vydrino Shoulder (51.58350°N, 104.85180°E; water depth 665 m; collected 1 Aug 2001).

The proposed treatment steps generally follow techniques used by Brown et al. (1989), Regnéll (1992), Regnéll and Everitt (1996), and Nakagawa et al. (1998), but include a modification in specific gravity and a final acid-alkali-acid treatment combined with 10  $\mu\text{m}$  micro-sieving. Heavy liquid separation combined with fine- and micro-sieving can provide pure sporomorph samples (Zhou et al. 1997; Morgenroth 2000). The steps for preparation of sporomorph concentrates are presented in Figure 2. The resulting sporomorph concentrate consists of grains of 10–90  $\mu\text{m}$  in diameter, with a density of 1.13–1.66  $\text{g}/\text{cm}^3$ .

### Sample Purity Control

The obtained concentrates were checked under the microscope and photographed (see Figure 3 for a photo example), and sample purity was estimated on the basis of particle counts. Pollen and spore spectra of the concentrates were generally similar to those obtained during standard pollen analysis. However, due to the additionally required chemical (hydrolysis, bleaching) and physical processing steps, some small and delicate pollen types (especially *Alnus fruticosa*) were partly lost. Organic matter, which could be removed from the sporomorph samples only in part, was comprised of lignin (conifer wood fragments), small globose limnic microfossils, and hyaline particles. For quantification of sample purity, the number of pollen and spores was related to the overall count of organic particles present in the concentrates. By estimation, the resulting purity was slightly adjusted according to the relative size of recorded non-sporomorph organic particles in comparison to pollen grain. Table 1 contains the estimated purity for all the samples. The best sample purity was obtained for the Vydrino core CON01-605-5 due to very high concentrations of large bisaccate pollen grains in the sediment.

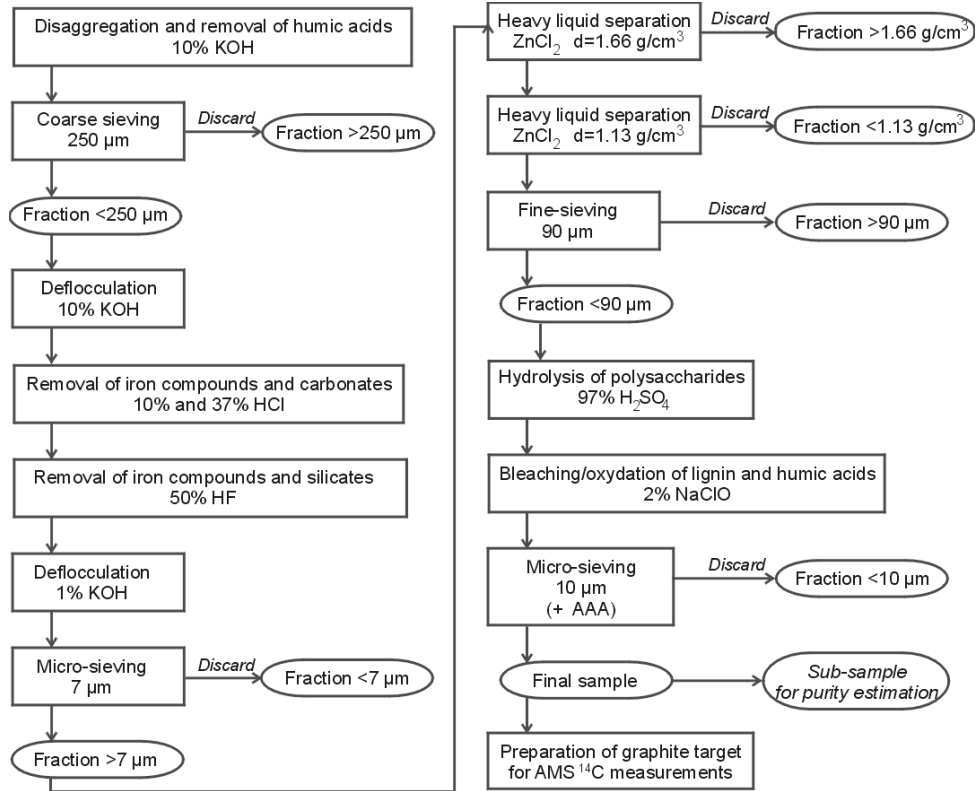


Figure 2 Flow-chart for preparation of sporomorph concentrates

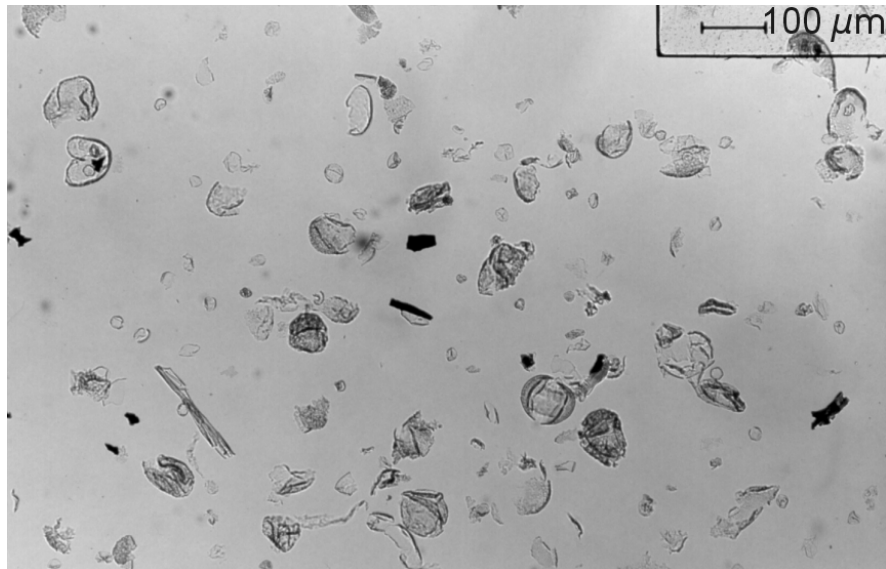


Figure 3 Photo documentation, sample LBF-AMS-8

### Preparation of Graphite Targets for AMS <sup>14</sup>C Dating

The extracts were prepared for AMS measurement in the Gliwice Radiocarbon Laboratory according to the procedures described in Goslar and Czernik (2000) and Czernik and Goslar (2001). The concentration of <sup>14</sup>C in produced graphite targets was measured in the Poznań Radiocarbon Laboratory (Goslar et al., forthcoming).

### RESULTS AND DISCUSSION

Details concerning sample preparation and results are given in Table 1. Most of the samples yielded CO<sub>2</sub> in amount corresponding to 1 or more mg of carbon, enabling the preparation of standard AMS targets. The average carbon content in pollen grains (about 64%) is in accordance with an assumed value (Erdtman 1969; Brown et al. 1989). Additionally, when the amount of CO<sub>2</sub> was sufficient, stable carbon isotope measurements were carried out in the Mass Spectrometry Laboratory of the Maria Curie-Skłodowska University in Lublin. The results of stable isotope measurements are within the range of values typical for terrestrial matter.

The <sup>14</sup>C ages of pollen samples were calibrated using the OxCal program (Bronk Ramsey 2001) and INTCAL98 calibration curve (Stuiver et al. 1998), except the last 3 dates from Posolskoe Bank (LBK). These dates were beyond the INTCAL98 curve and calibration was performed using the CalPal program and CalPal 2001 curve (Weninger et al. 2003; Jöris and Weninger 1998). For the construction of timescales, the ranges of calibrated ages obtained for 2 confidence intervals were taken, and the centers of the probability distribution ranges were taken as point-estimates for linear regression calculations. For 2 dates of the Posolskoe Bank (LBK-AMS-5a and b), we could have narrowed down the obtained ranges by applying the calibration of an ordered sequence of <sup>14</sup>C dates (using OxCal), taking into consideration the stratigraphic order of samples. A similar procedure could be applied to samples LBK-AMS-8 to 10, but it seems not justified at this time because the INTCAL98 calibration curve does not cover this time interval. However, we believe that the new officially accepted curve would help to solve this problem.

The results obtained for Continent Ridge and Vydrino Shoulder show a high linear correlation and were used for estimating the average sedimentation rates (Figure 4a,b). The ages of Posolskoe Bank samples are more scattered, indicating the disturbances in the sedimentation process. Two periods of sedimentation were distinguished for which the average sedimentation rates were estimated (Figure 4c). The sedimentological studies, which are still in process, will enable the recognition of possible sedimentation disturbances.

For all 3 sites, a shift in obtained sediment ages can be noticed, resulting in a non-zero age for surface sediment. Various reasons for this effect can be considered. The most likely reason is an incomplete recovery of the core (i.e. with the topmost layer missing), but the other ones also have to be considered. The surface age effect seems to be a complex problem caused by more than 1 reason. A similar effect was observed in previous studies of Lake Baikal sediments (Coleman et al. 1996; Horiuchi et al. 2000), using AMS <sup>14</sup>C dating carried out on the total organic carbon (TOC) fraction. Extraction of pollen grains was undertaken in the hope of eliminating this effect, and because macrofossils are extremely rare in Lake Baikal sediments. However, the effect of inconsistent older ages obtained on pollen extracts was observed in other studies, such as the Lake Gościąg sediments (Kilian et al. 2002), which may lead to the conclusion that this could result from specific properties of pollen grains. The surface age effect will be investigated in further studies, and the dating of more samples from the uppermost core sections and from material suspended in water is planned. We hope that this will enable the creation of reliable timescales for the sedimentation process.

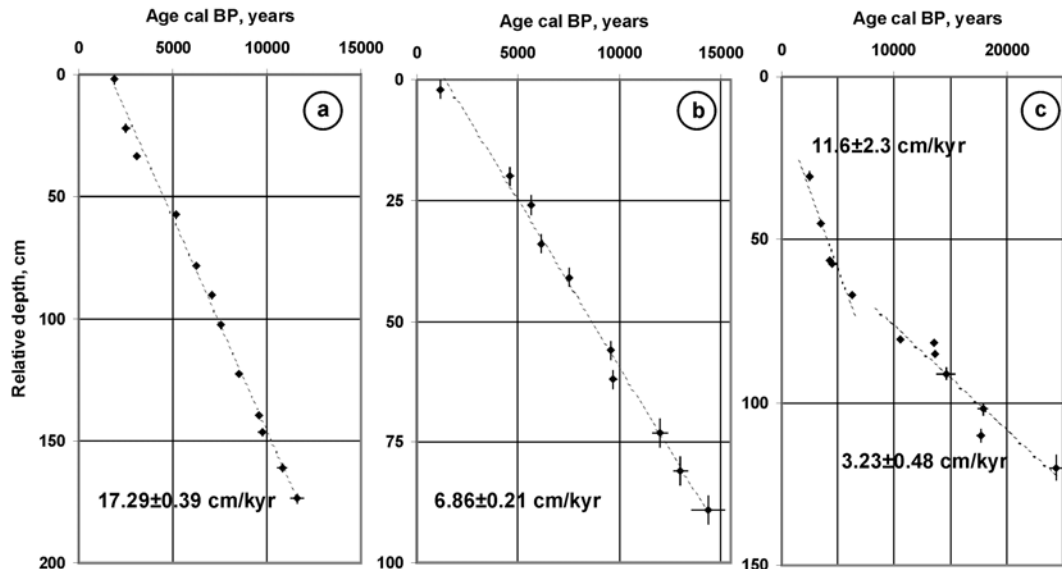


Figure 4 Calibrated age of samples versus depth for 3 investigated cores (a–Vydrino, b–Continent, c–Posolskoe). The vertical error bars represent the depth range and horizontal bars represent the range of calibrated ages. The average sedimentation rates based on linear regression are given in the graphs.

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